A photograph of a desert landscape with a dirt road leading towards a volcano in the distance. The sky is blue with some clouds. The text is overlaid on the image.

# Formation of iron oxide and phosphate tephra of the El Laco Volcano, Chile from an Fe-P-O-S magma

James Mungall

Kristy Long

James Brennan

Richard Naslund

# Outline

- Define and describe El Laco/Kiruna type Fe oxide deposits; demonstrate that they are magmatic in origin
  - Large scale structures
  - Field occurrences
  - Petrography, mineral compositions, fluid inclusions
  - Occurrence of coexisting silicate and oxide melt in same samples
- Discuss petrogenetic processes – how can oxide magmas form?
- New experimental observations bearing on liquid immiscibility
- Discussion of phase relations
- Problems and solutions to petrogenetic riddles

# El Laco-type Fe ore deposits

- Worldwide distribution but rare
- Proterozoic to Tertiary ages
- Massive magnetite and hematite
  - Low  $\text{TiO}_2$ , low  $\text{SiO}_2$
  - Up to a few % apatite
  - Trace amounts of clinopyroxene, quartz
  - REE rich
  - Spatial association with hydrothermal IOCG deposits
  - Associated with potassic arc magma suites
- NOT nelsonites (i.e., Ti-rich magnetite-apatite bodies)
- NOT ilmenite deposits like Tellnes or Tio deposits
- Igneous textures and structures
  - Massive dikes
  - Vesicular and brecciated flows
  - Intrusive breccias
  - Unconsolidated ash and crystal tuffs
- Common association with intense hydrothermal alteration
- Supposed by some workers to be hydrothermal themselves

# Composition of the Ore (without minor veins)

$\Sigma \text{FeO} \sim 98 \%$

$\text{MgO} = 0.5 - 1.3 \%$

$\text{SiO}_2 = 0.0 - 1.0\%$

$\text{P}_2\text{O}_5 = 0.0 - 1.0 \%$

$\text{TiO}_2 = 0.15 - 0.85 \%$

$\text{REE} = 5000-10000$

$\text{V} = 900 - 2500 \text{ ppm}$

$\text{Ni} = 90 - 360 \text{ ppm}$

$\text{Mn} = 300 - 700 \text{ ppm}$

$\text{Zr} = 10 - 100 \text{ ppm}$

$\text{Al Ca Na K} \lll 1\%$



# El Laco

- Central stratovolcano dominated by two-pyroxene biotite-bearing andesite ( $>16\text{km}^3$ ) and late central rhyodacite
- Late adventive lava flows composed of magnetite and hematite
- $2.0 \pm 0.3 \text{ Ma}$  (Gardweg & Ramirez, 1985)
- Five main magnetite-hematite ore bodies with reserves =  $\sim 500 \text{ Mt}$  ( $\sim 0.1 \text{ km}^3$ )
- Dikes & veins of magnetite, pyroxene, and apatite in ore and host rock ( $<1\%$  of ore)
- Abundant pervasive alteration of host andesite & fumarolic sulfate mounds.

San Vicente Bajo

Rodados Negros

San  
Vicente  
Alto

El Lago volcano

Pico Laco  
(5470 m)

Laco Norte

Laco Sur

Image © 2009 DigitalGlobe  
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© 2009 Google

Imagery Date: May 9, 2006

23°49'57.14" S 67°29'13.54" W elev 4776 m

Eye alt 12.32 km



May 22, 2009 9:11 am

San Vicente Bajo

San  
Vicente  
Alto

Rodados Negros

El Laco volcano

Pico Laco  
(5470 m)

Laco Norte

Laco Sur

1 km

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Imagery Date: May 9, 2006

23°50'48.72" S 67°29'24.66" W elev 4525 m

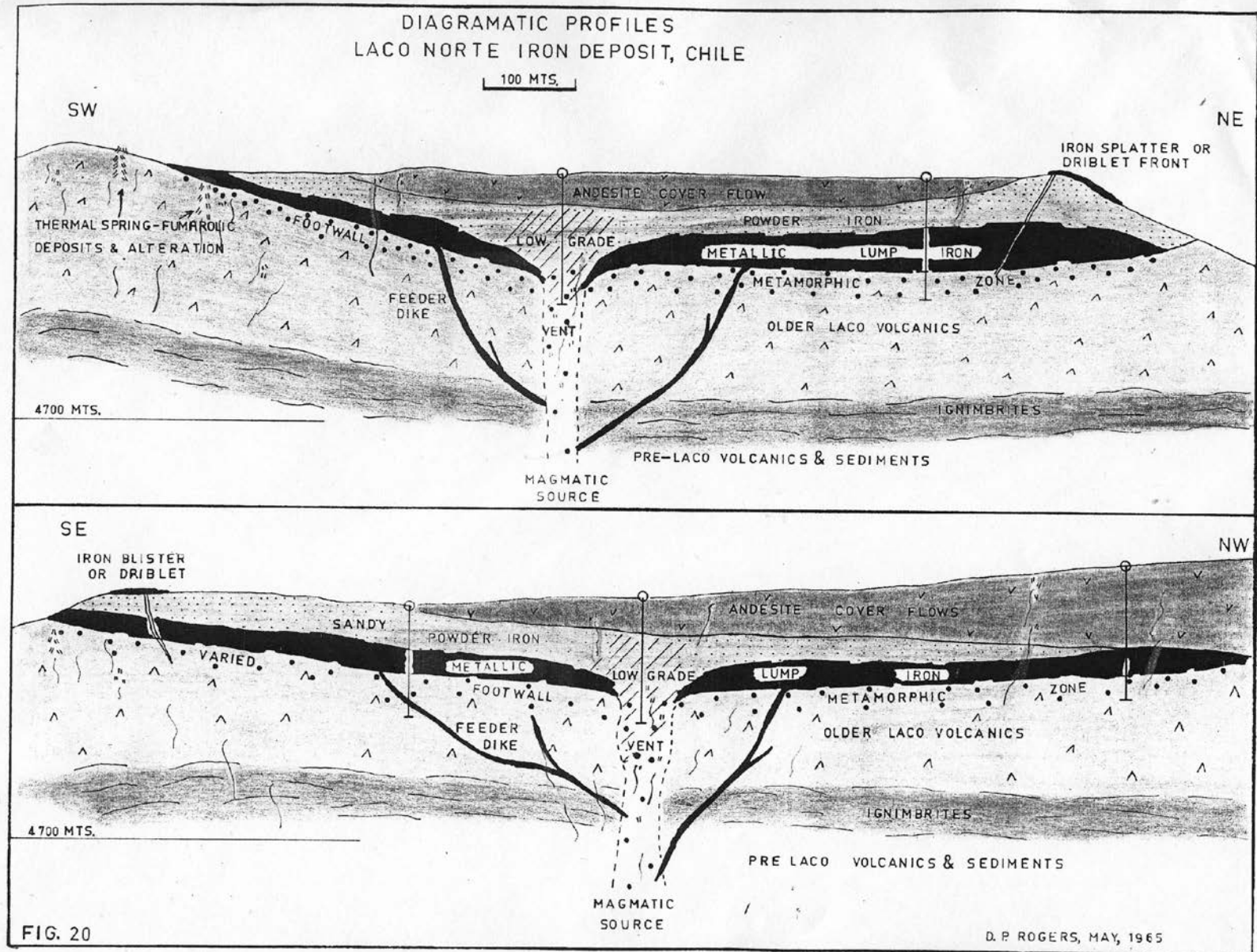
Eye alt 5.77 km

# Magnetite Lava Flow El Laco

↑  
~100m  
↓

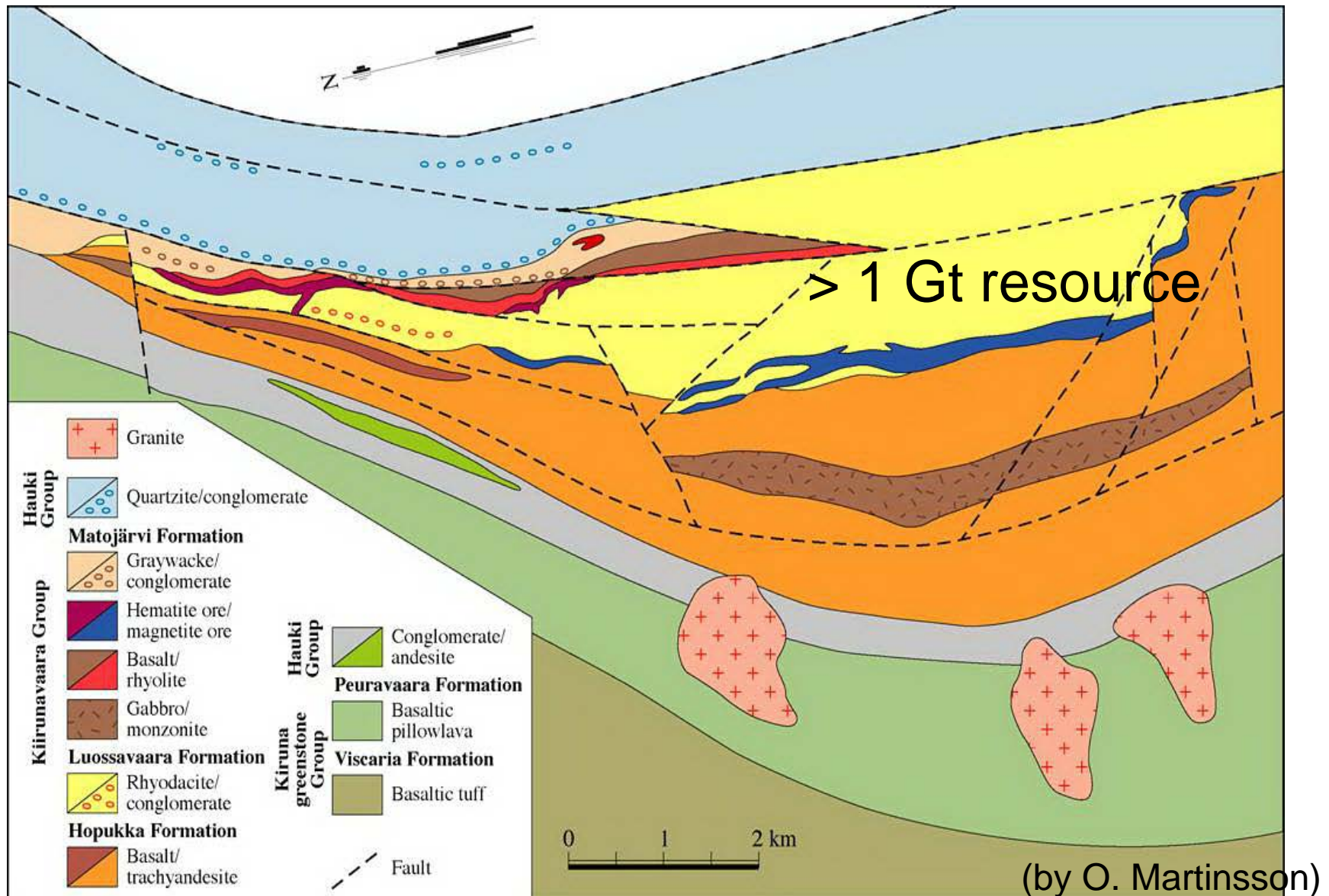
Photo: HR Naslund

# Map scale

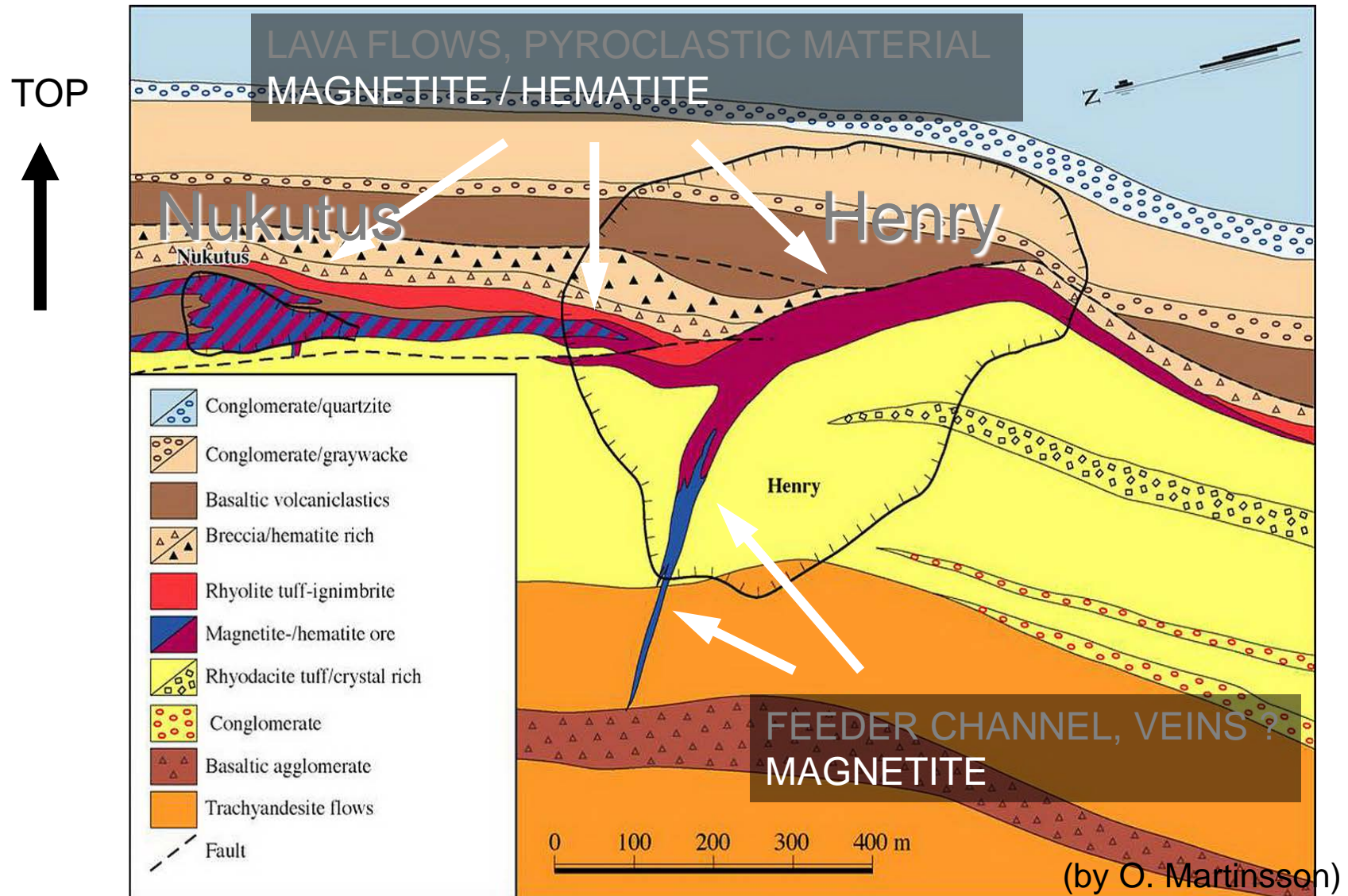




# Geology in the Kiruna area



# STRATIGRAPHY OF HENRY AND NUKUTUS, Kiruna district, Sweden



# Map scale observations

- Massive magnetite dikes, sills
- Gradation into magnetite(hematite) flows
- Hematite(magnetite) pyroclastic deposits



# Outcrop scale observations



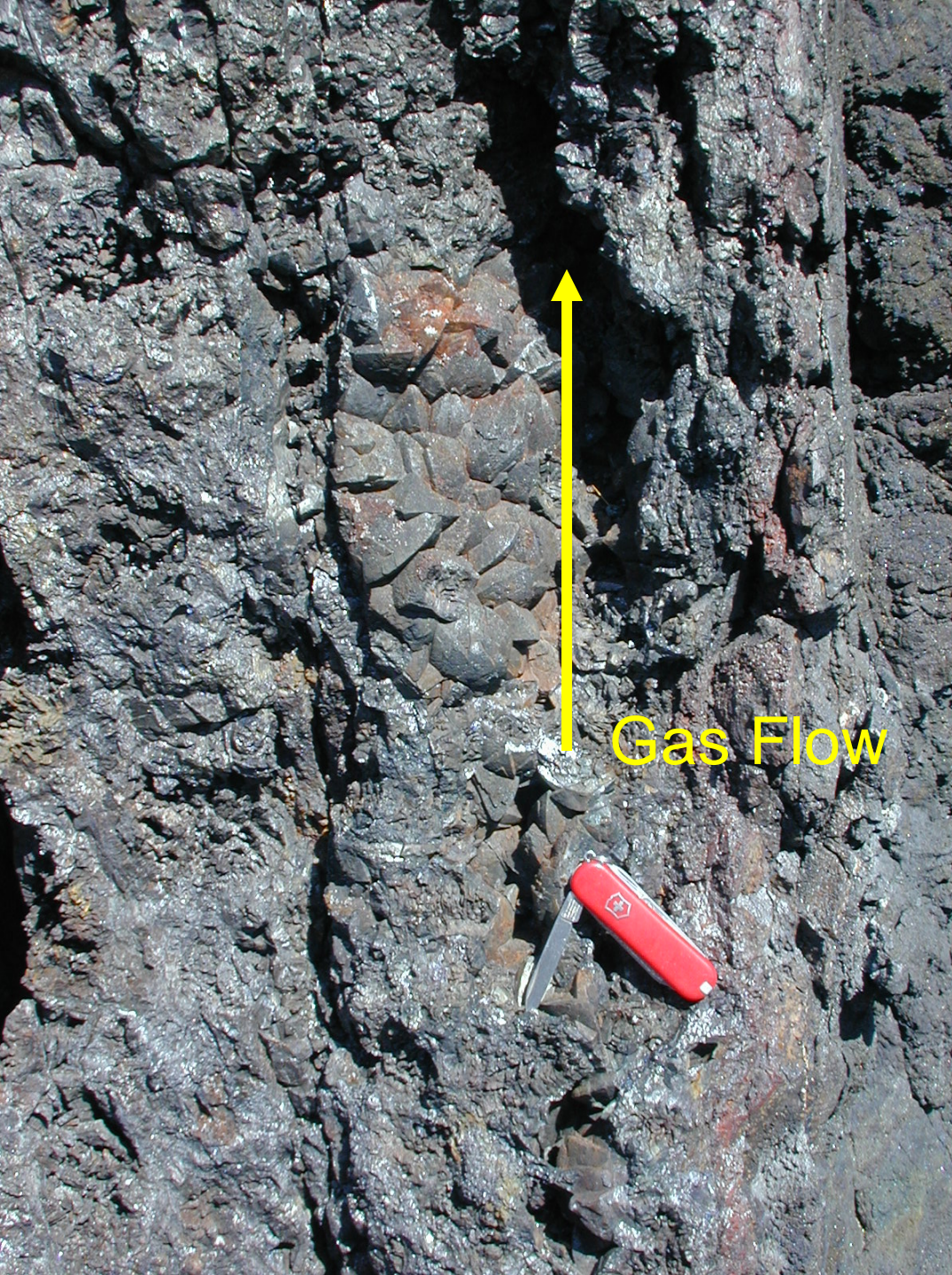
<-- 6 cm -->

## Vesicular ore at Laco Sur

Typical Ore Densities 1.9 - 3.1

Photo: HR Naslund





**Gas Escape Tube  
Laco Sur**

**Lined with  
Magnetite**

# Kirunavaara

## Vesicular iron ore





# Ore contact at Laco Sur



Razor sharp  
contact  
between iron  
ore and  
altered, iron-  
poor  
andesite.



Altered Andesite

Fine-grained  
Contact zone--  
Chilled Margin

Coarse  
porous  
iron ore



Photo: HR Naslund



**Magnetite dike  
w/ sharp contacts  
cutting host rock  
below Laco Sur**

**Chilled  
Margin**



Photo: HR Naslund



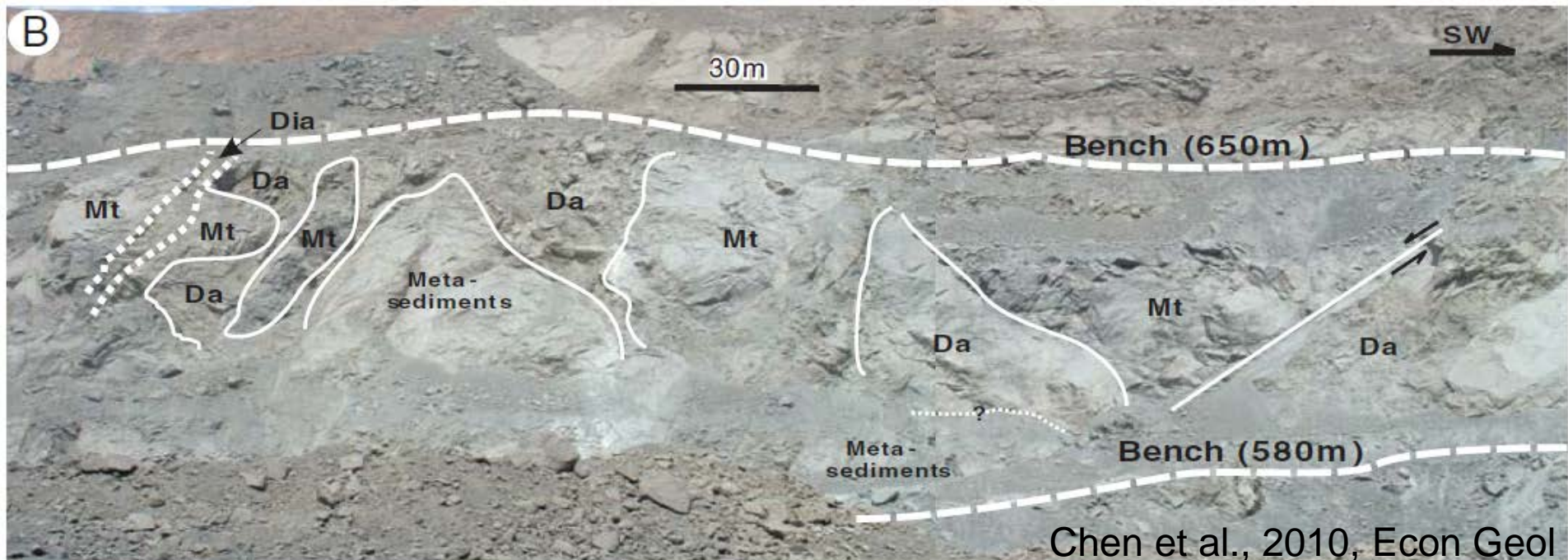
# Outcrop scale



- Kirunavaara
- Ore breccia in the footwall



# Marcona 1.9 Gt (Peru)







Magnetite Scoria  
Laco Norte

Photo: HR Naslund





Magnetite volcanic  
bomb(?)  
in hematite ash

Photo: HR Naslund



Diadochite =  $\text{Fe}_2(\text{PO}_4)(\text{SO}_4)(\text{OH}) \cdot 5\text{H}_2\text{O}$



Photo: HR Naslund

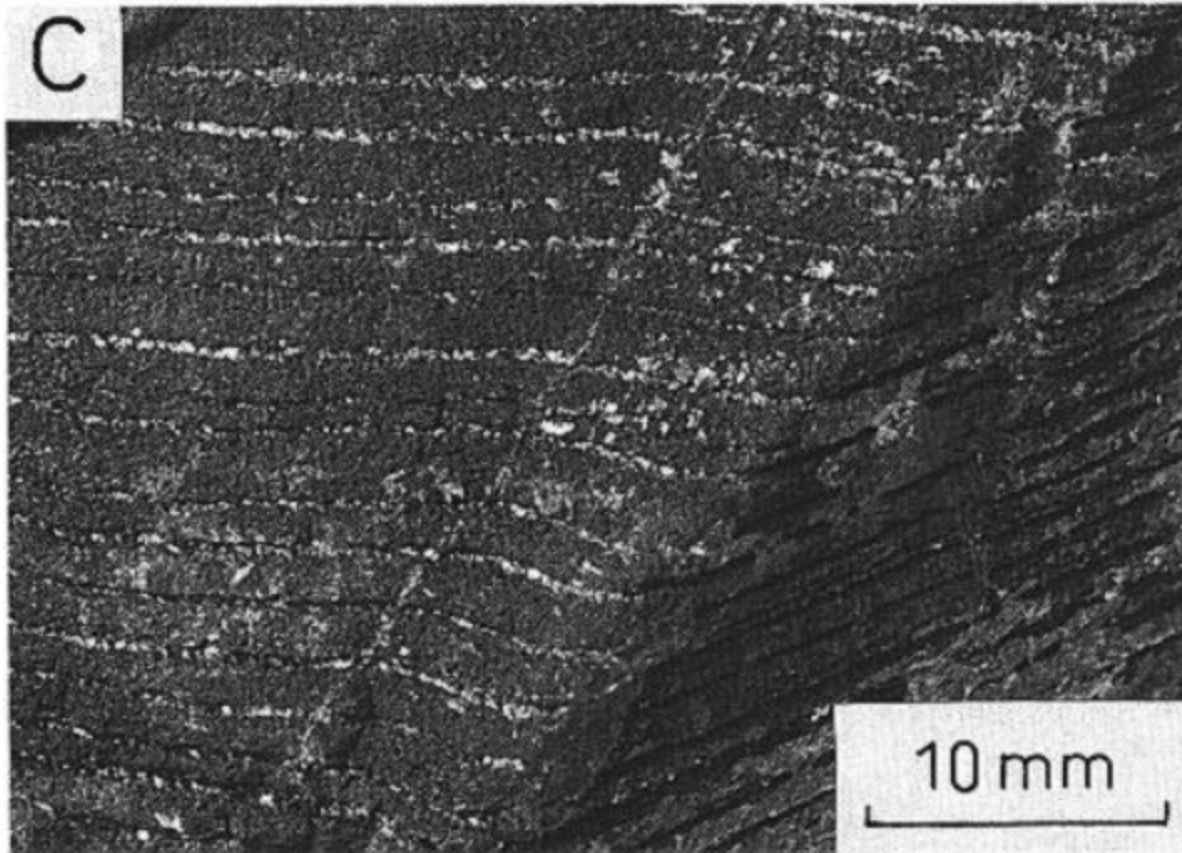


Platey magnetite, Laco Sur tephra; Nystrom and Henriquez, Econ Geol, 1994



# Kirunavaara

- Crystal tuff
- Magnetite, apatite



# Durango, Mexico

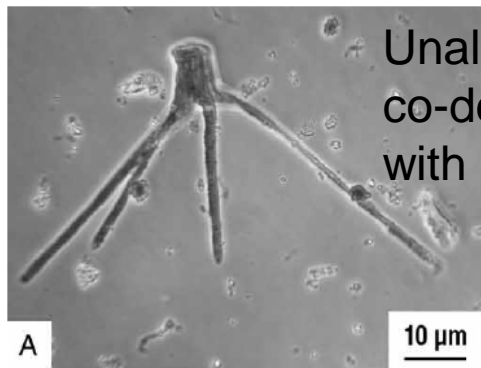


Iron oxide  
ignimbrite

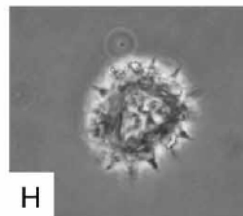
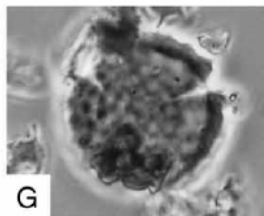
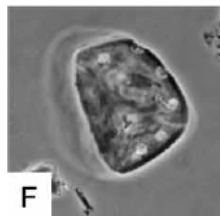
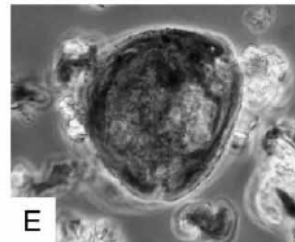
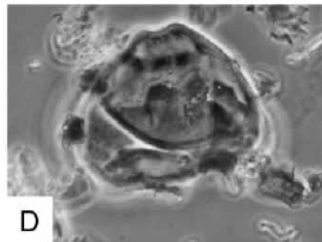
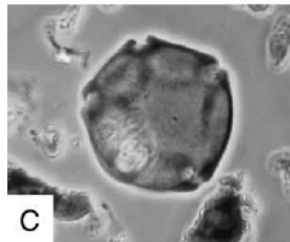
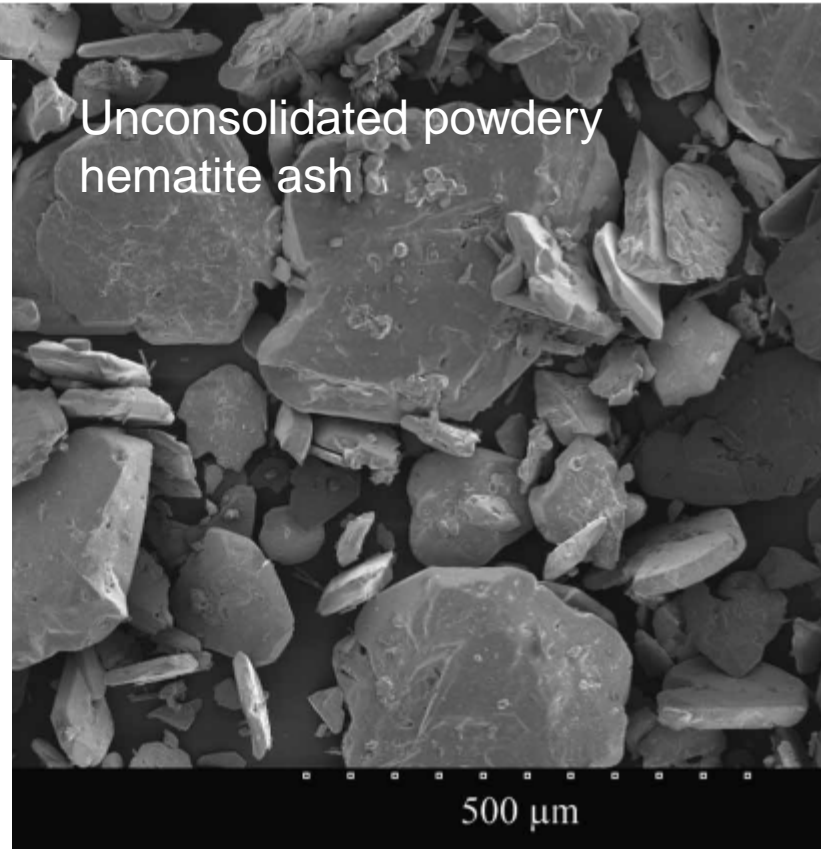


Vesicular iron ore

# La Perla, Mexico



Unaltered, unheated pollen  
co-deposited  
with ash



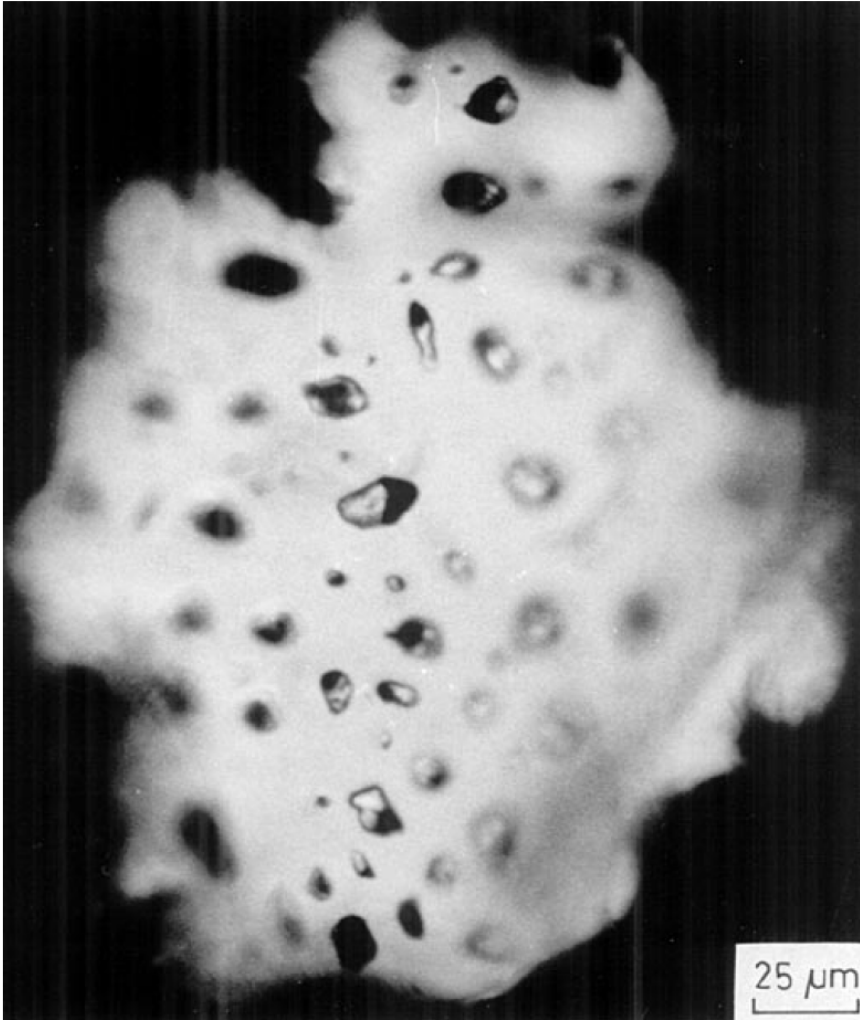


# Outcrop scale observations

- Perfect preservation of complex primary magmatic structures
  - Vesicles
  - Unconsolidated air-fall crystal tuffs
  - Preservation of unaltered (cold) pollen grains
  - Chilled margins on dikes, sills
- Magnetite breccias, dikes, sills, flows
- Hematite in pyroclastic materials
- Intrusive breccias
- Plutonic zones of mingling of oxide and silicate magma
- Locally intense but not universal hydrothermal alteration



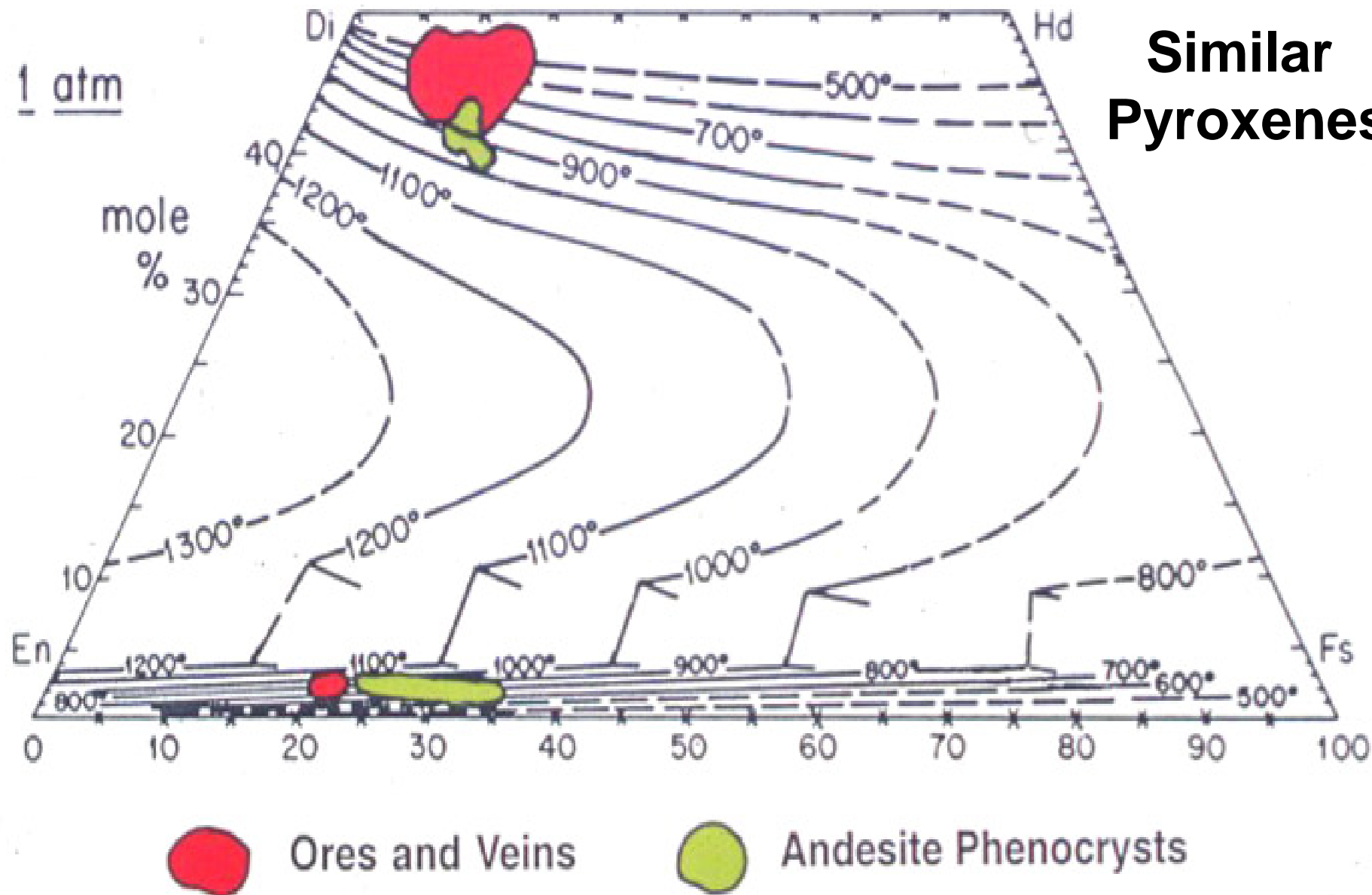
# Micro-scale observations at El Laco

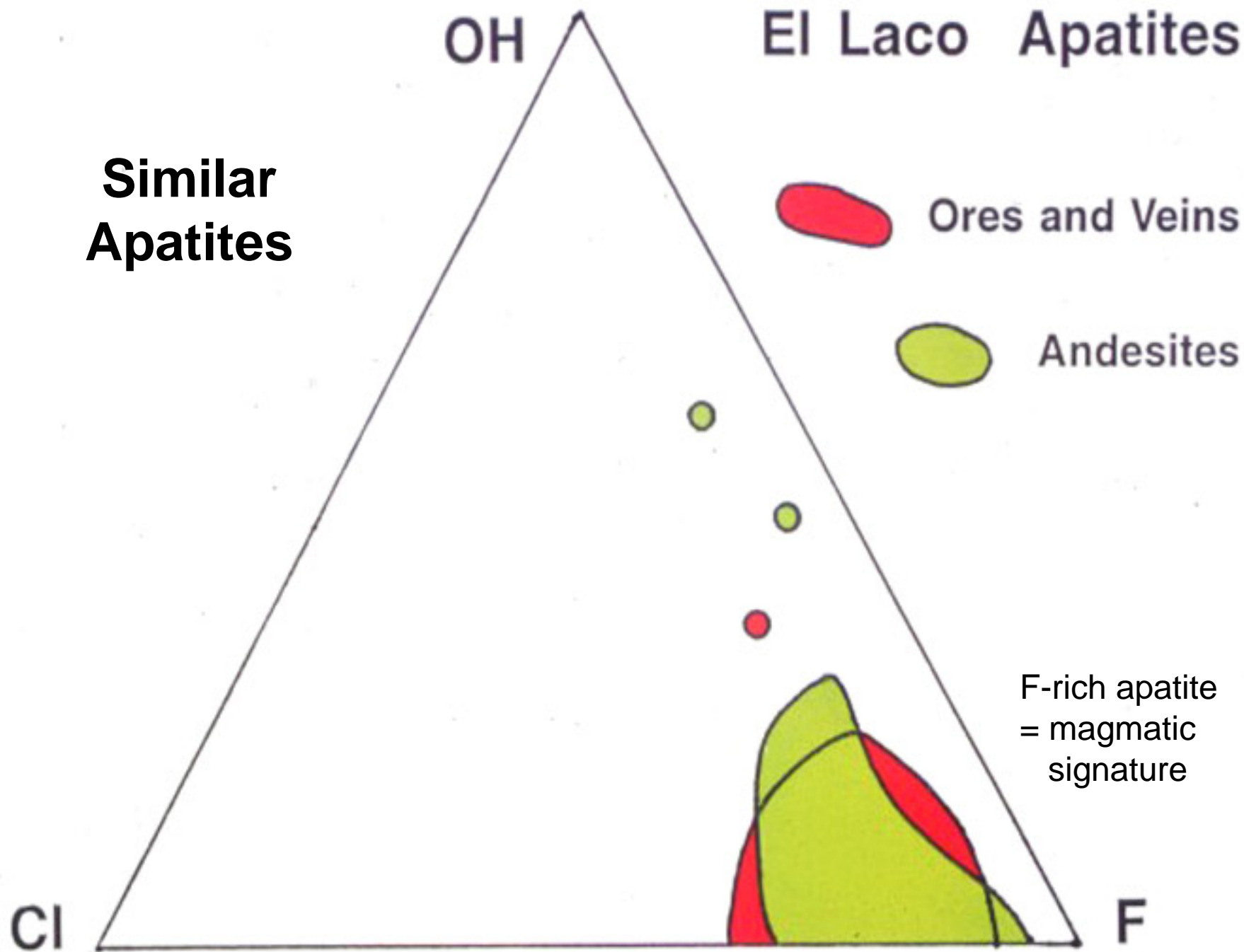


- Hydrous saline melt inclusions in pyroxene and apatite
- Homogenization temperatures > 800°C

LINDSLEY: PYROXENE THERMOMETRY

**Similar  
Pyroxenes**





# New petrographic observations from El Laco

- Tephra cooled rapidly, hence may preserve traces of quenched igneous phases
- Petrography of oxide-phosphate tephra
- Preliminary melting experiments
- Bulk composition of ash:
  - 86 mol%  $\text{Fe}_2\text{O}_3$
  - 12 mol%  $\text{FePO}_4$
  - 2 mol%  $\text{SiO}_2$

El Laco  
ash



hematite

monazite



20 kV

300X

50  $\mu\text{m}$

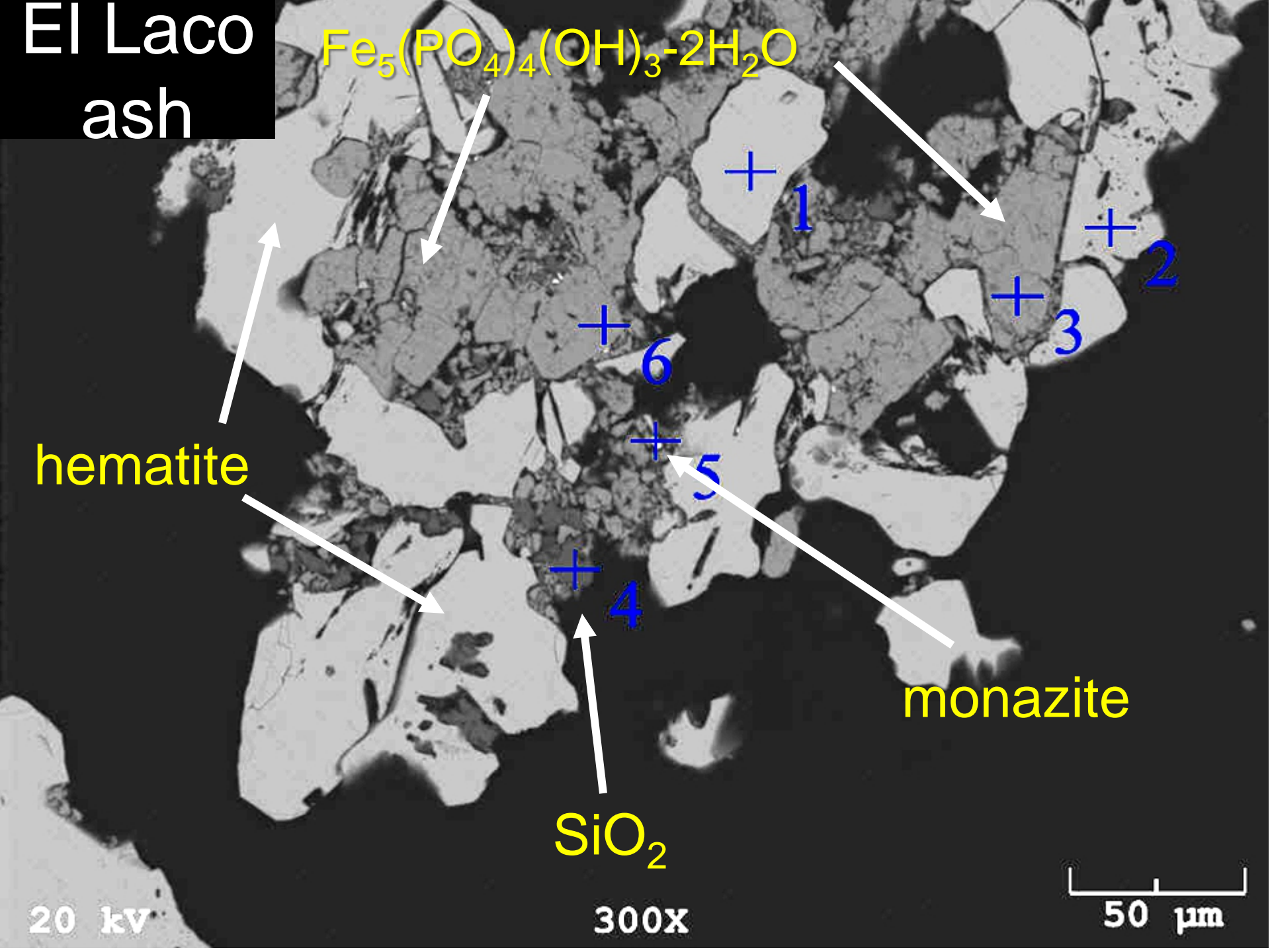


Image22-3

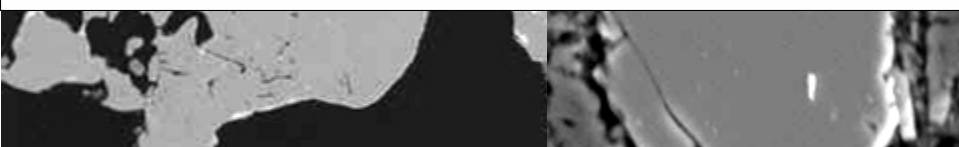
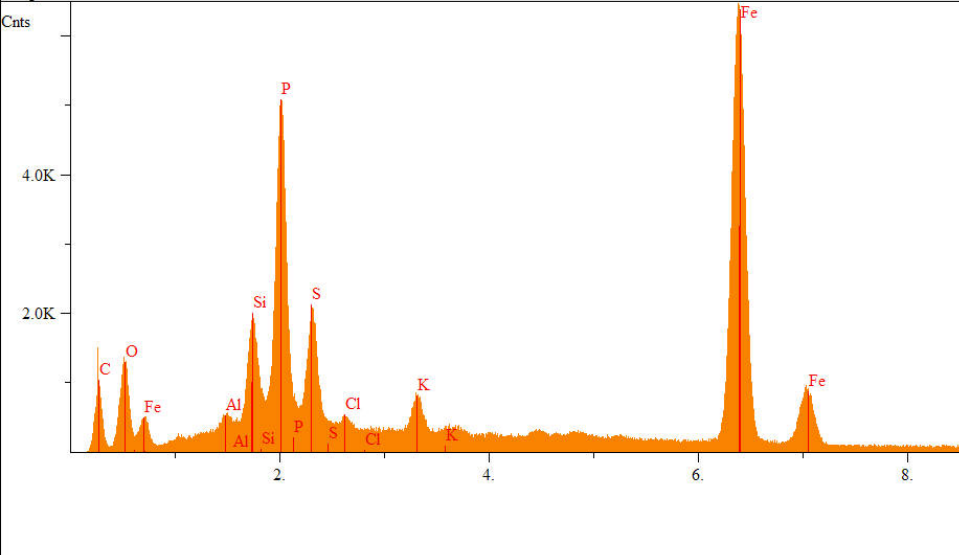
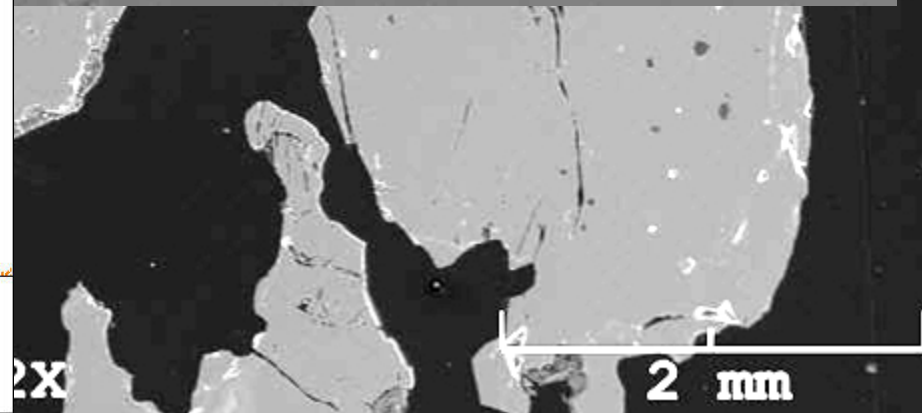
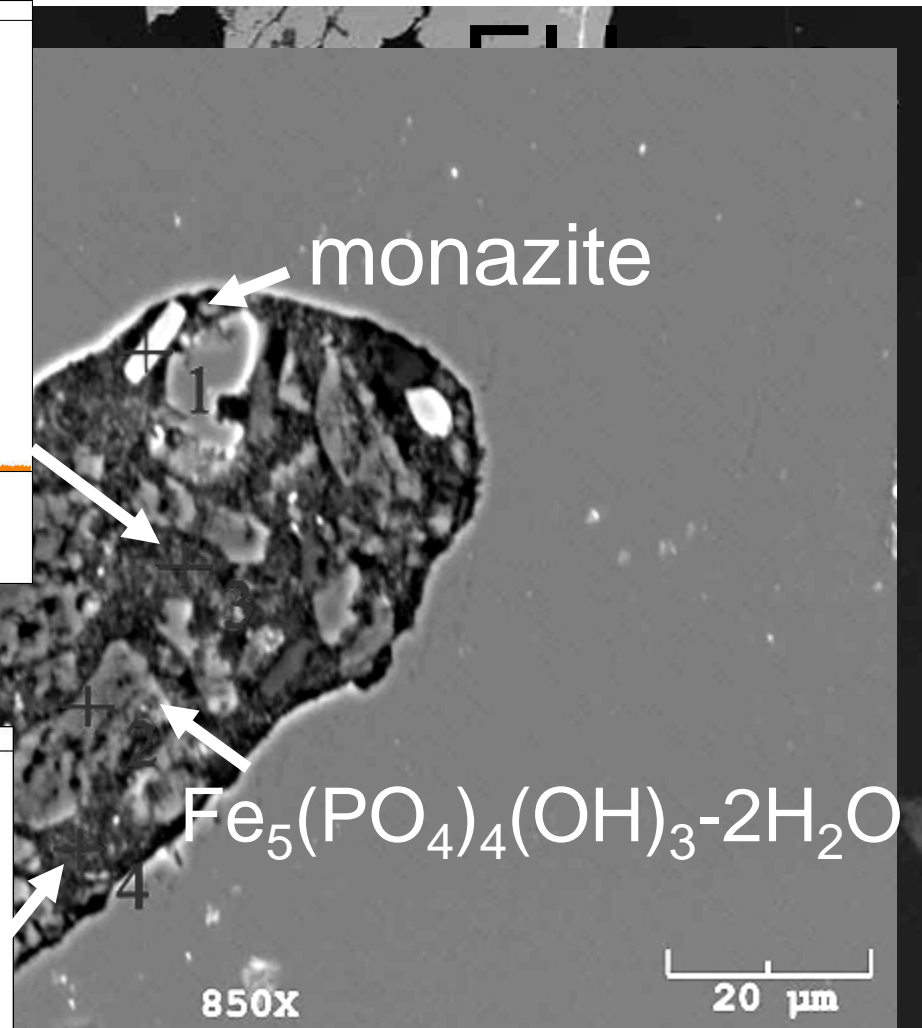
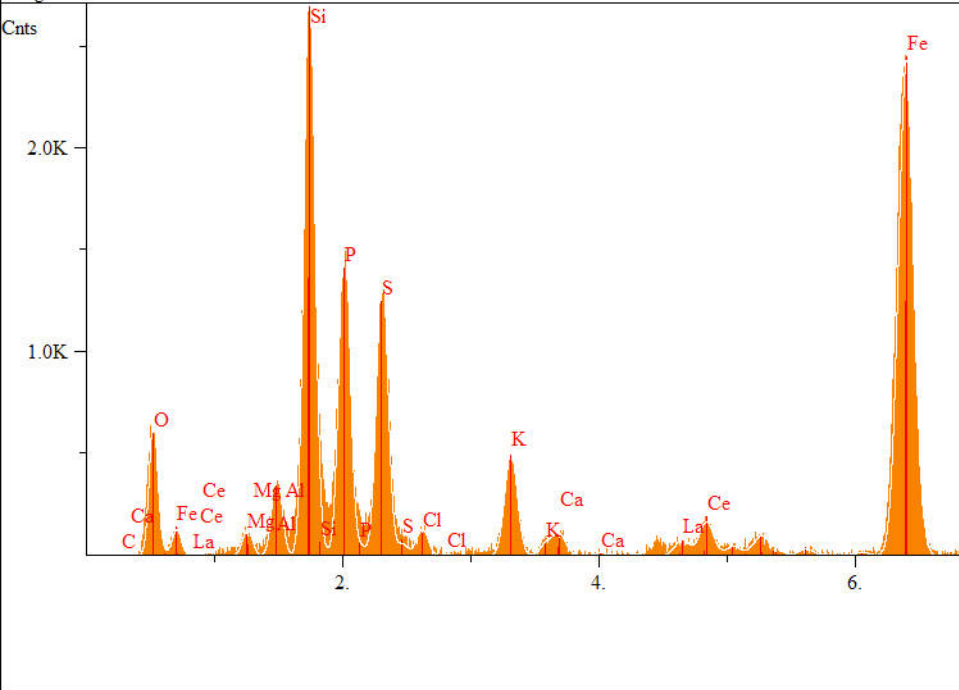


Image22-4

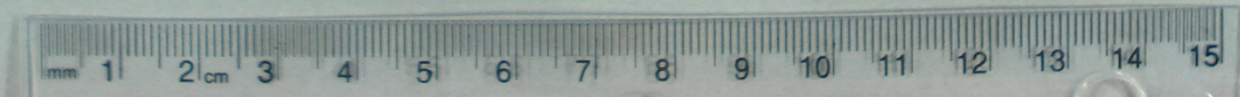


# El Laco ash observations

- Predominantly hematite, which armors patches of other phases
- Fe-P phases important – identified elsewhere at El Laco by Haggerty (1968) as  $\text{Fe}_3^{2+}\text{Fe}^{3+}(\text{PO}_4)_3$  (anhydrous)
- Accessory Ca-rich monazite and  $\text{SiO}_2$
- Subsidiary S, Cl (wt%)
- S presumably as  $\text{SO}_3$  in presence of Mt, Hm

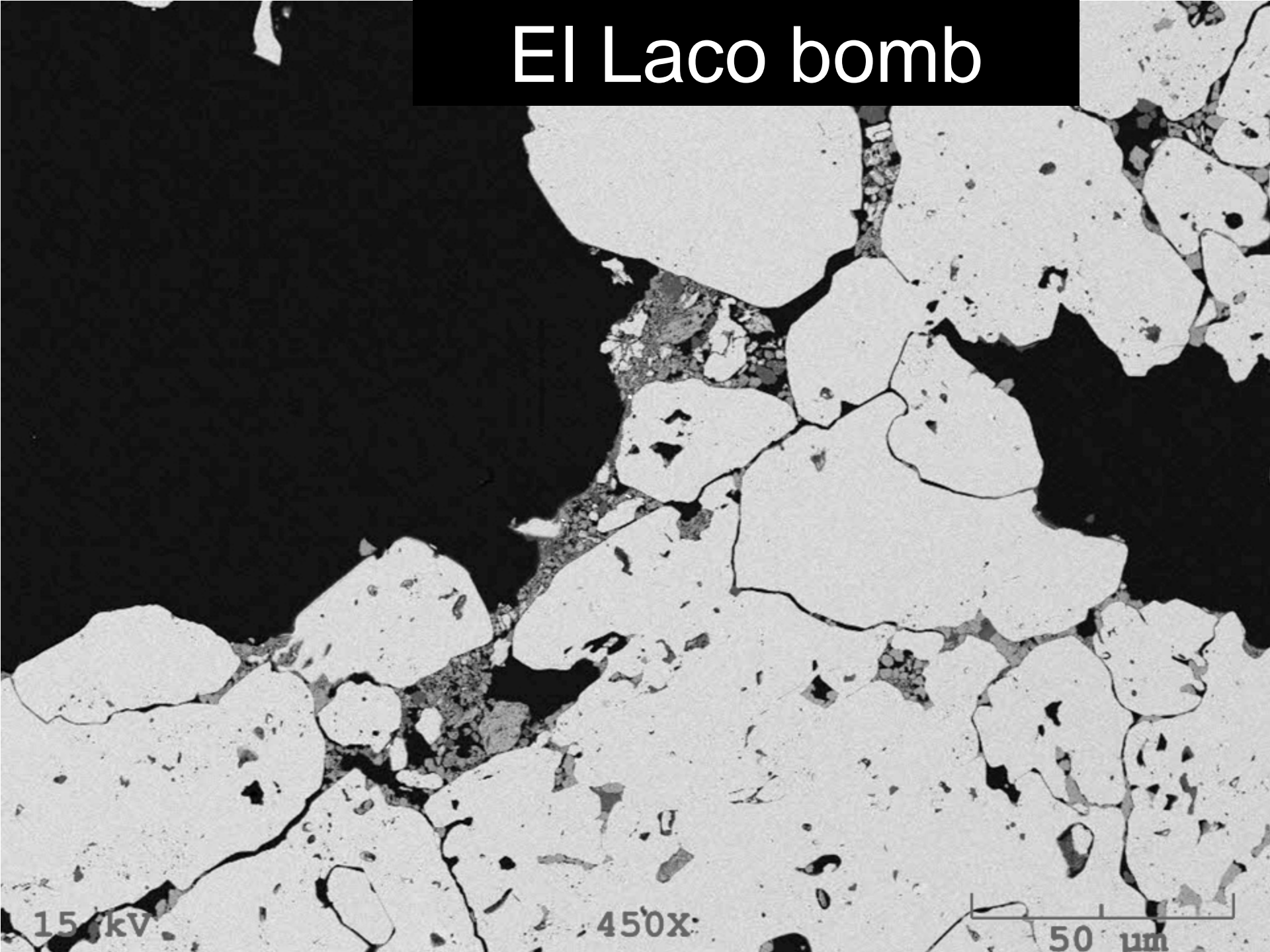


# El Laco bomb (El Bob)





# El Lago bomb

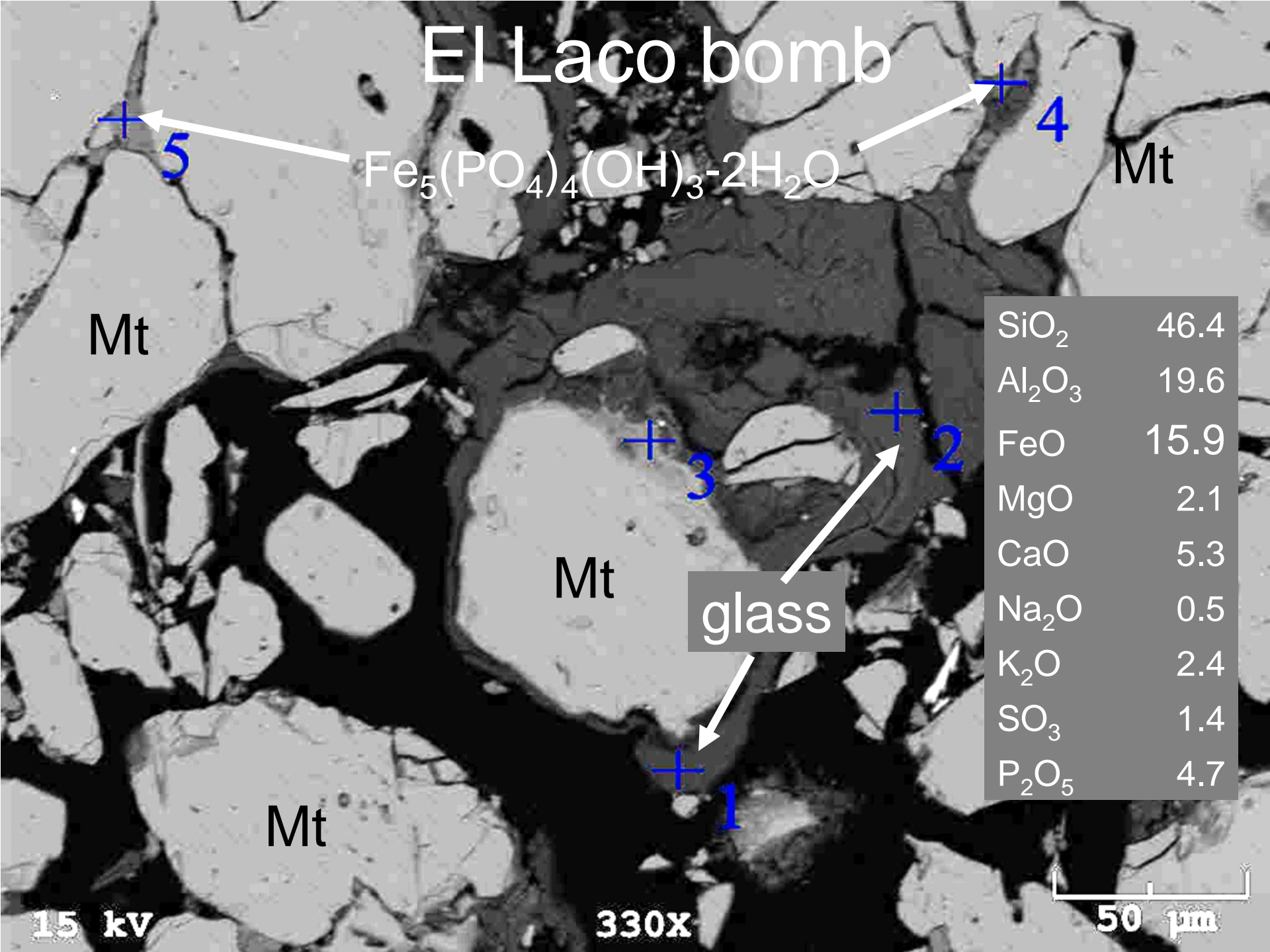


15 kV

450X

50 μm

# El Laco bomb



SiO <sub>2</sub>	46.4
Al <sub>2</sub> O <sub>3</sub>	19.6
FeO	15.9
MgO	2.1
CaO	5.3
Na <sub>2</sub> O	0.5
K <sub>2</sub> O	2.4
SO <sub>3</sub>	1.4
P <sub>2</sub> O <sub>5</sub>	4.7

# El Laco bomb



Mt

glass

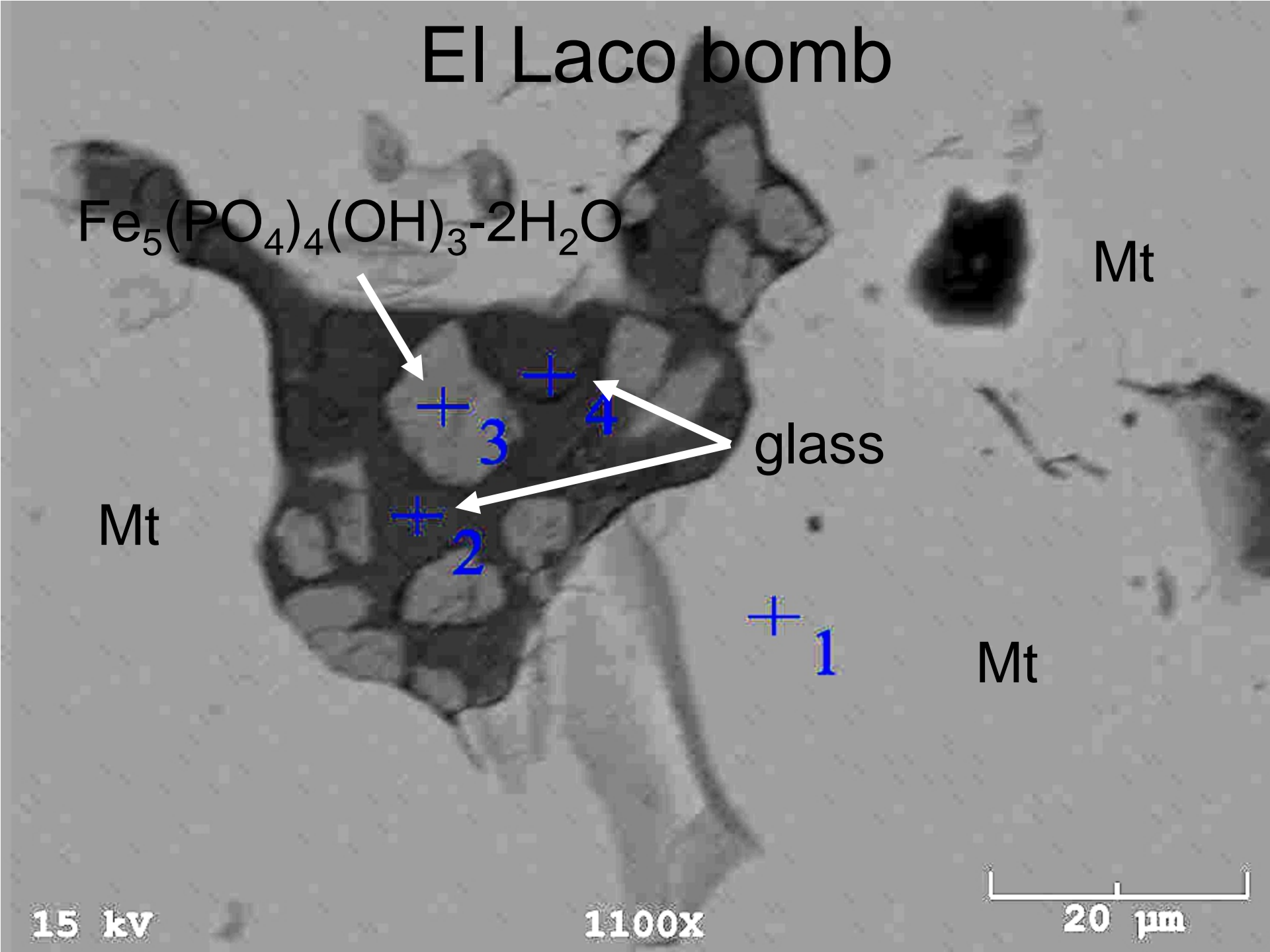
Mt

Mt

15 kV

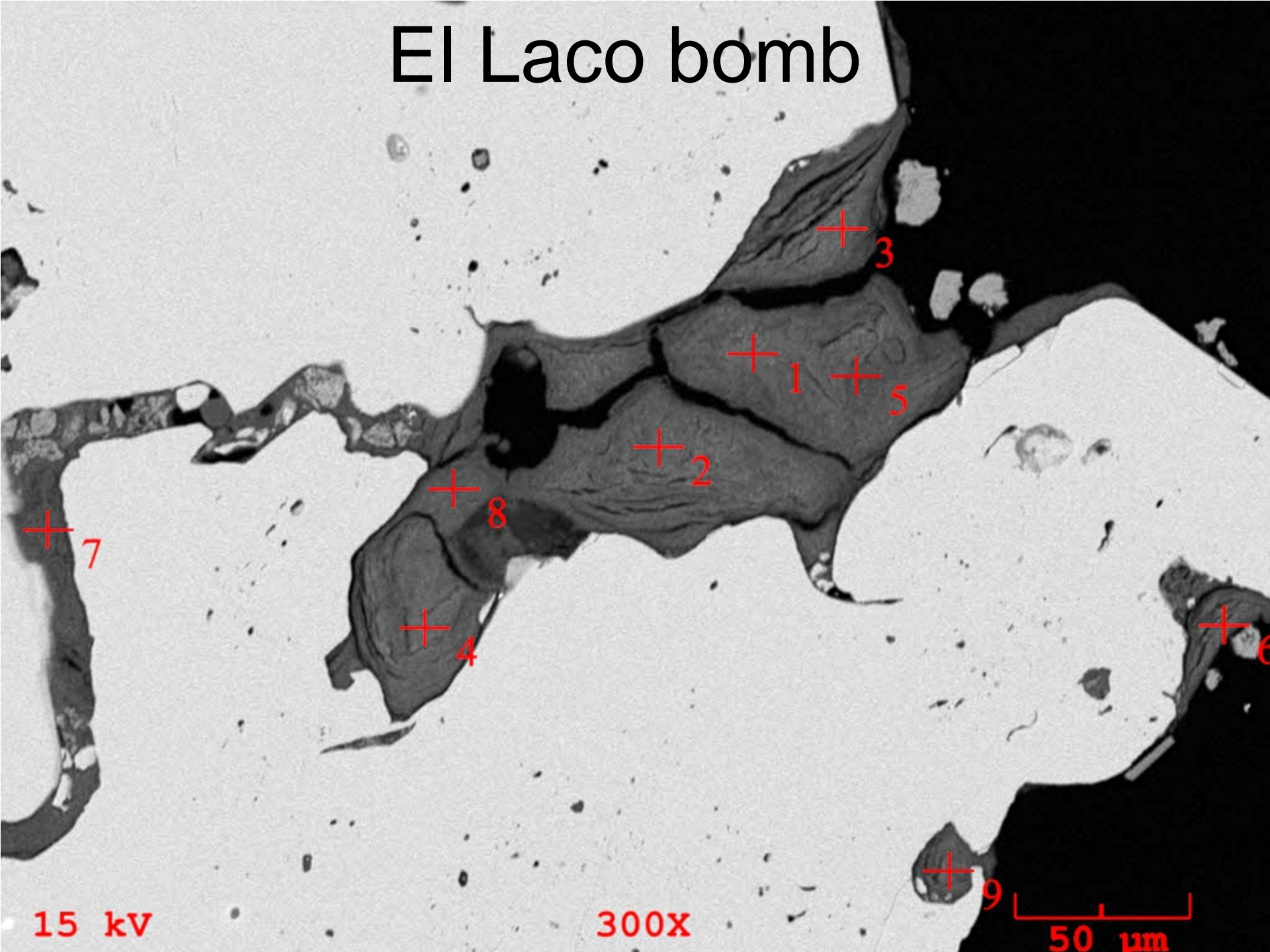
1100x

20  $\mu\text{m}$





# El Lago bomb



15 kV

300X

50 μm

# El Laco bomb observations

- Similar mineral assemblage as ash but more magnetite than hematite
- Similar intratelluric phosphate assemblage
- Also includes films of potassic basaltic glass intermittently coating interior surfaces
- Glass appears to be in contact (textural equilibrium?) with all solid phases including Fe phosphates
- Both glass and Fe phosphate accessory phases appear to form a meniscus between hematite crystals

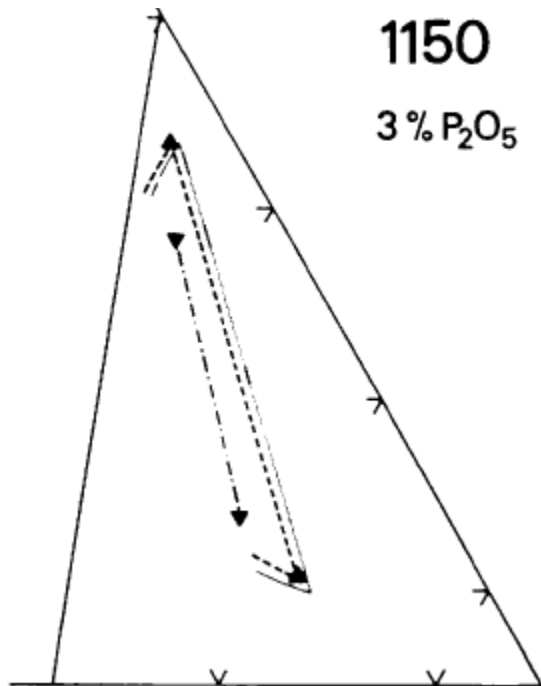
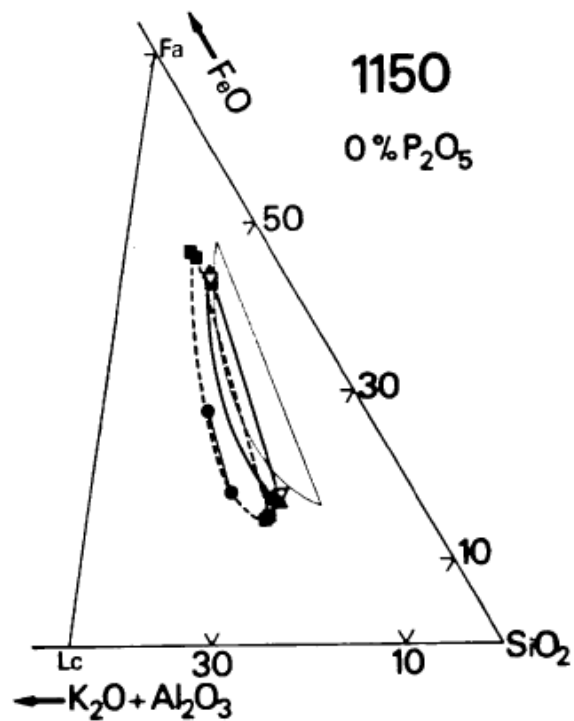
# Genesis of Fe oxide magmas

- Partial melting
  - What on Earth could melt to form this composition? Nothing.
- Fractional crystallization
  - No known silicate magma fractionates to  $\text{Fe}_2\text{O}_3$  enrichment and  $\text{SiO}_2$  exhaustion
- Liquid immiscibility
  - Iron rich silicate melts commonly show immiscible phase separation in the laboratory
  - Numerous field occurrences on microscopic scale of immiscible Fe-rich and Si-rich liquids
  - Problem: these examples are all too  $\text{SiO}_2$ -rich to correspond to magnetite magmas



# Experimental occurrence of liquid immiscibility

- $\text{K}_2\text{O}-\text{Al}_2\text{O}_3-\text{FeO}-\text{SiO}_2$  system
  - Naslund, 1983, Am Mineral
  - Visser and Koster van Groos, 1979a,b, Am J Sci
  - Watson, 1976, CMP
- $\text{K}_2\text{O}-\text{Al}_2\text{O}_3-\text{FeO}-\text{SiO}_2$  system
  - Roedder 1978, GCA
- Magnetite-apatite-"diorite"
  - Philpots 1967, Econ Geol



# Leucite-fayalite-silica

- Immiscible liquids
- Works just as well in sodic magmas
- Too much silica for Fe-oxide magma

Photo: HR Naslund

Visser & Koster van Groos, 1979a,b, Am J Sci



# Liquid immiscibility: predictions for the field

- Cosaturation of both liquids with the same phases
- Continued exsolution of liquids during cooling
- Textural evidence for coexistence of two melts
- “The best type of evidence for the existence of one liquid as an immiscible separation from another, and one that would involve the element of interpretation in minimum degree, would be the occurrence of glassy globules in a glassy rock of different composition”

N. L. Bowen, 1928

The evolution of the Igneous Rocks

# Mesoscopic liquid immiscibility

- Mertainen, Sweden
- Trachyte host rocks
- Magnetite-rich globules
- Ore veins and breccias

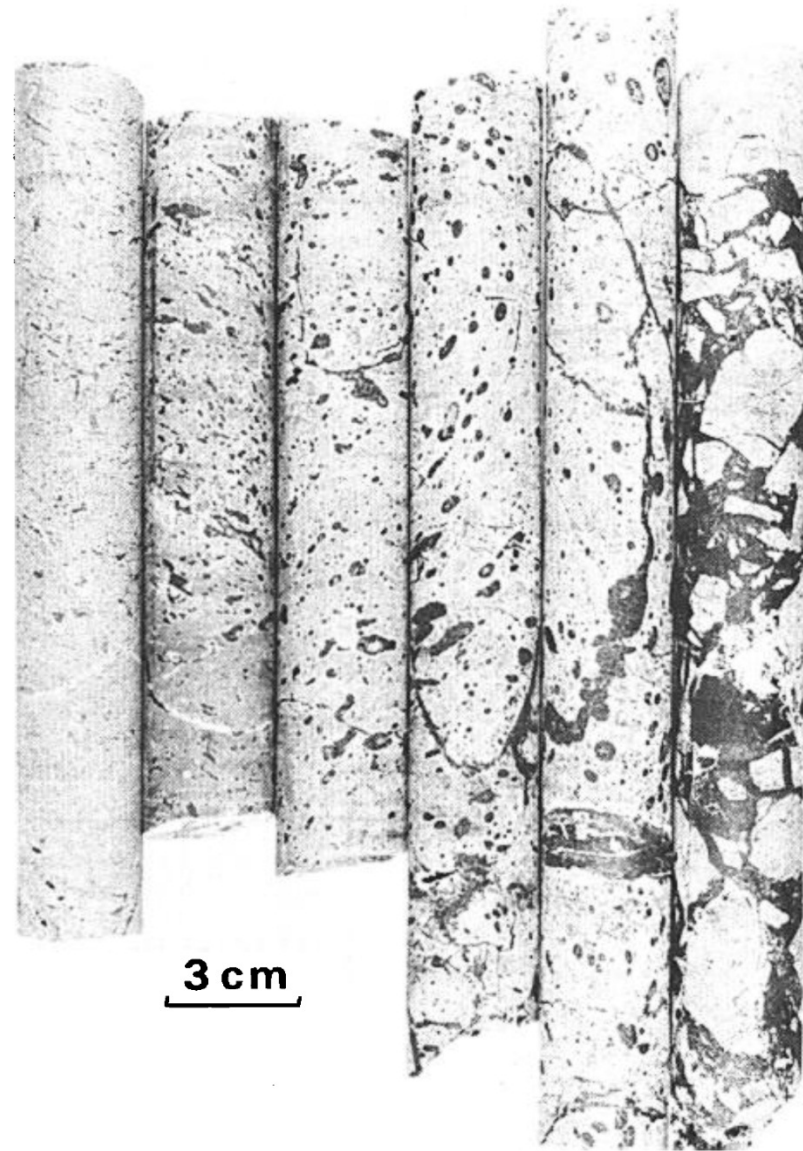


FIG. 16. Core samples from the Mertainen area (DH 63303). From left to right in the photograph are illustrated the textural forms of magnetite within the host trachyte. The textures primarily show the variation in size and shape of the magnetite-rich globular shapes, some of which are connected by a series of magnetite-infilled fractures or veins. Note the lighter colored depletion zones around the globules and fractures within the trachyte host rock.

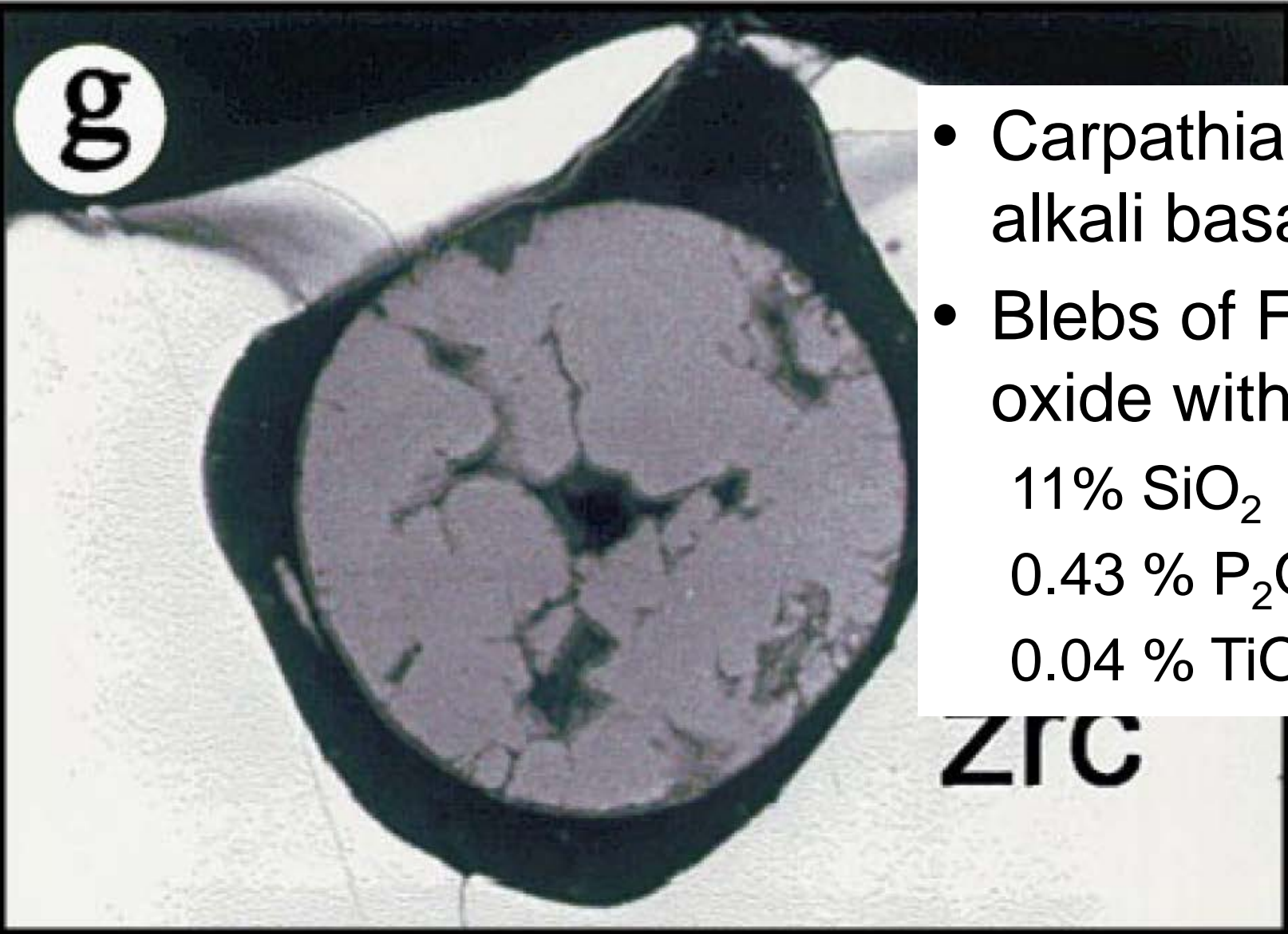


# Lunar basalts



- Metastable liquid immiscibility
- Below liquidus of supercooled magma

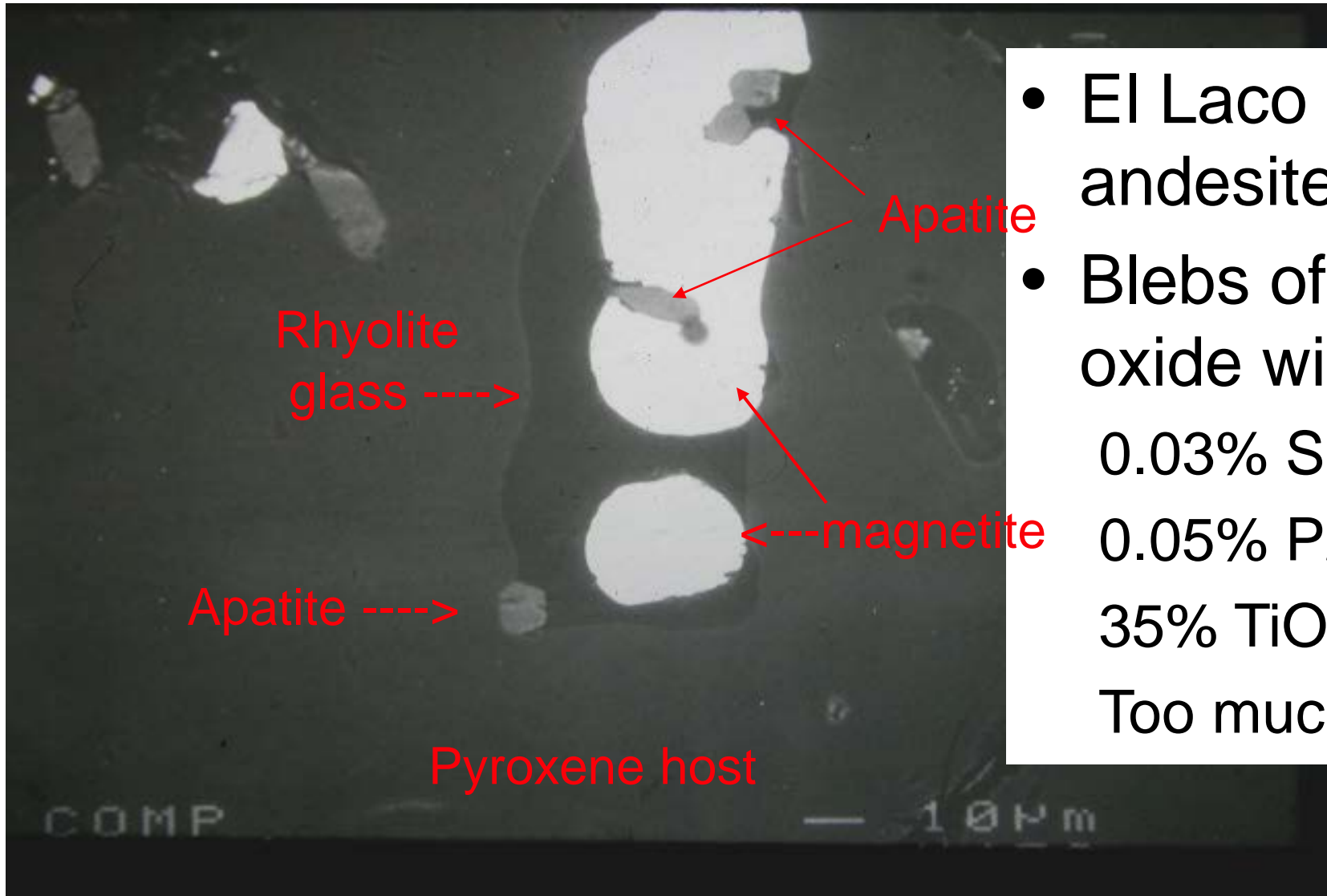
# Melt inclusions:



- Carpathian alkali basalt
- Blebs of Fe oxide with  
11%  $\text{SiO}_2$   
0.43 %  $\text{P}_2\text{O}_5$   
0.04 %  $\text{TiO}_2$

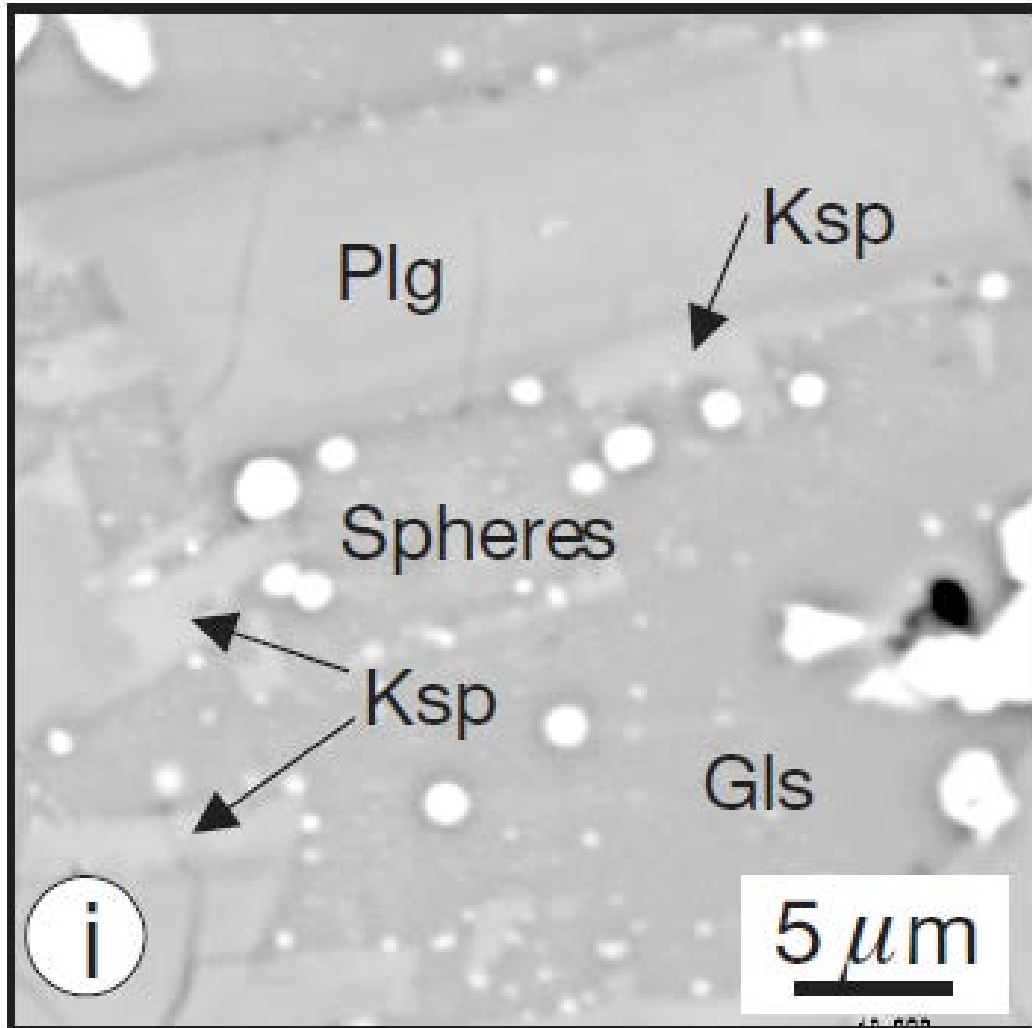


# Melt inclusions



- El Laco andesite
- Blebs of Fe oxide with  
0.03%  $\text{SiO}_2$   
0.05%  $\text{P}_2\text{O}_5$   
35%  $\text{TiO}_2$   
Too much Ti

# Antauta volcanic centre, Peru

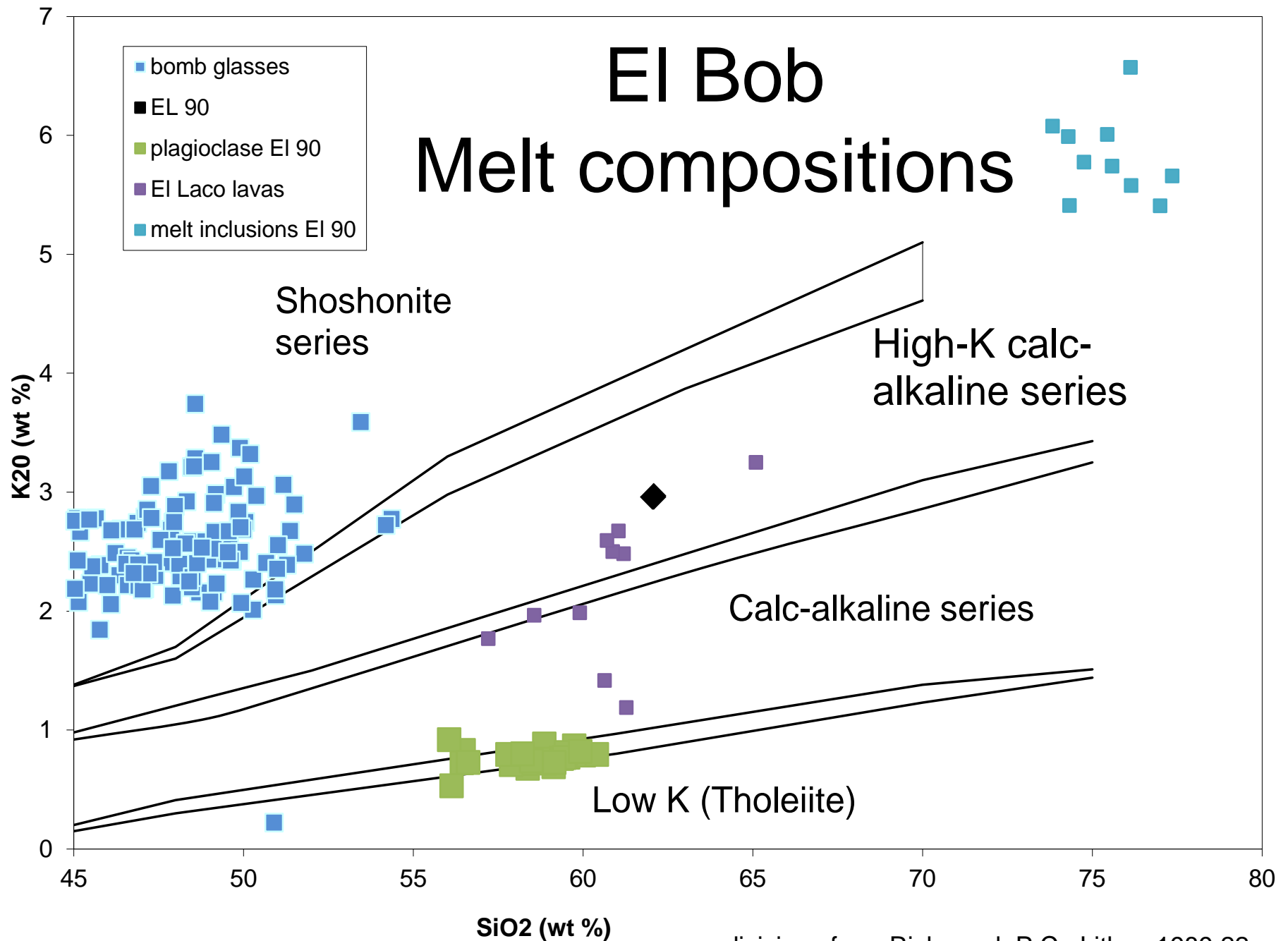
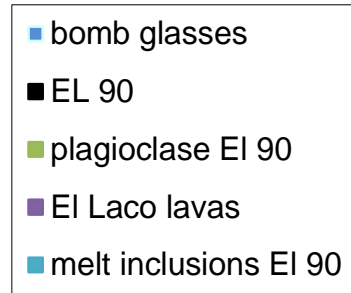


- Immiscible Fe-Ti rich liquids
- Absarokite – rhyodacite magma mixing
- Fe oxide with
  - ≥ 8.4 % SiO<sub>2</sub>
  - ≥ 23 % P<sub>2</sub>O<sub>5</sub>
  - ≥ 30 % TiO<sub>2</sub>
- Too much Ti

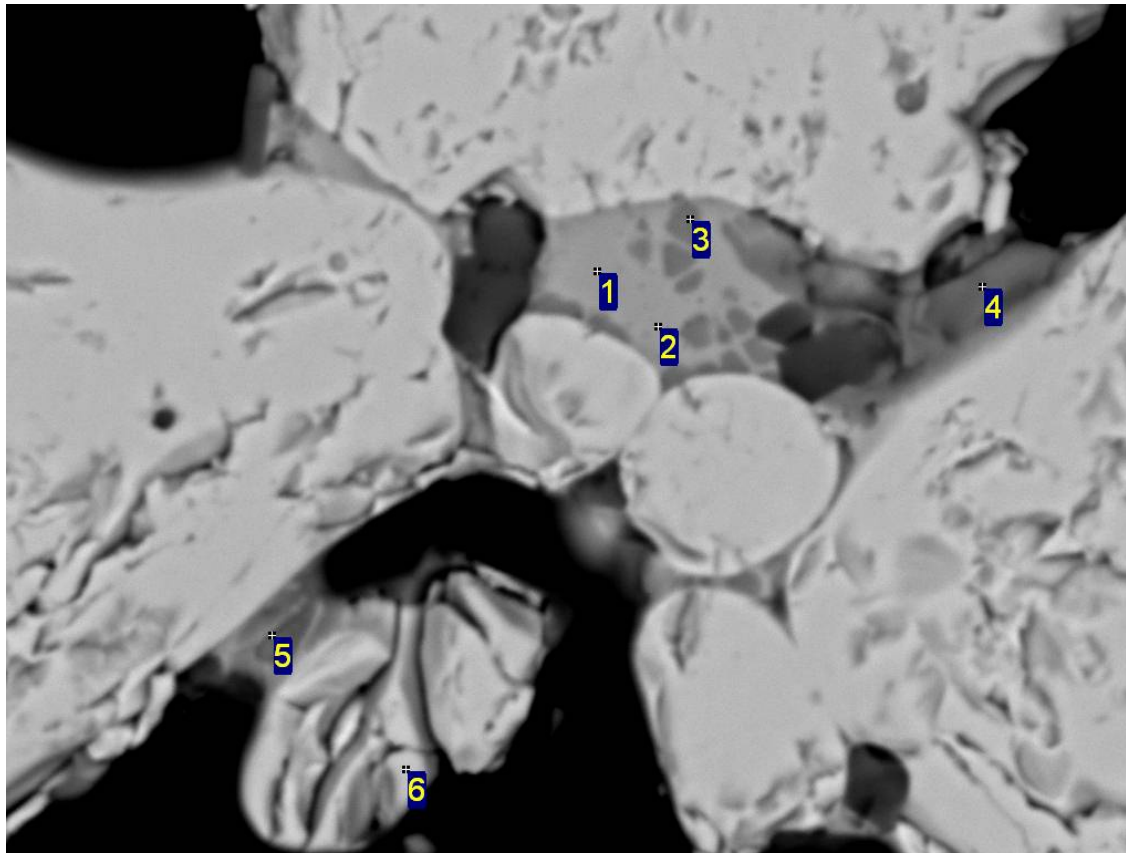


# El Bob

## Melt compositions



# 1 Atm melting of ash

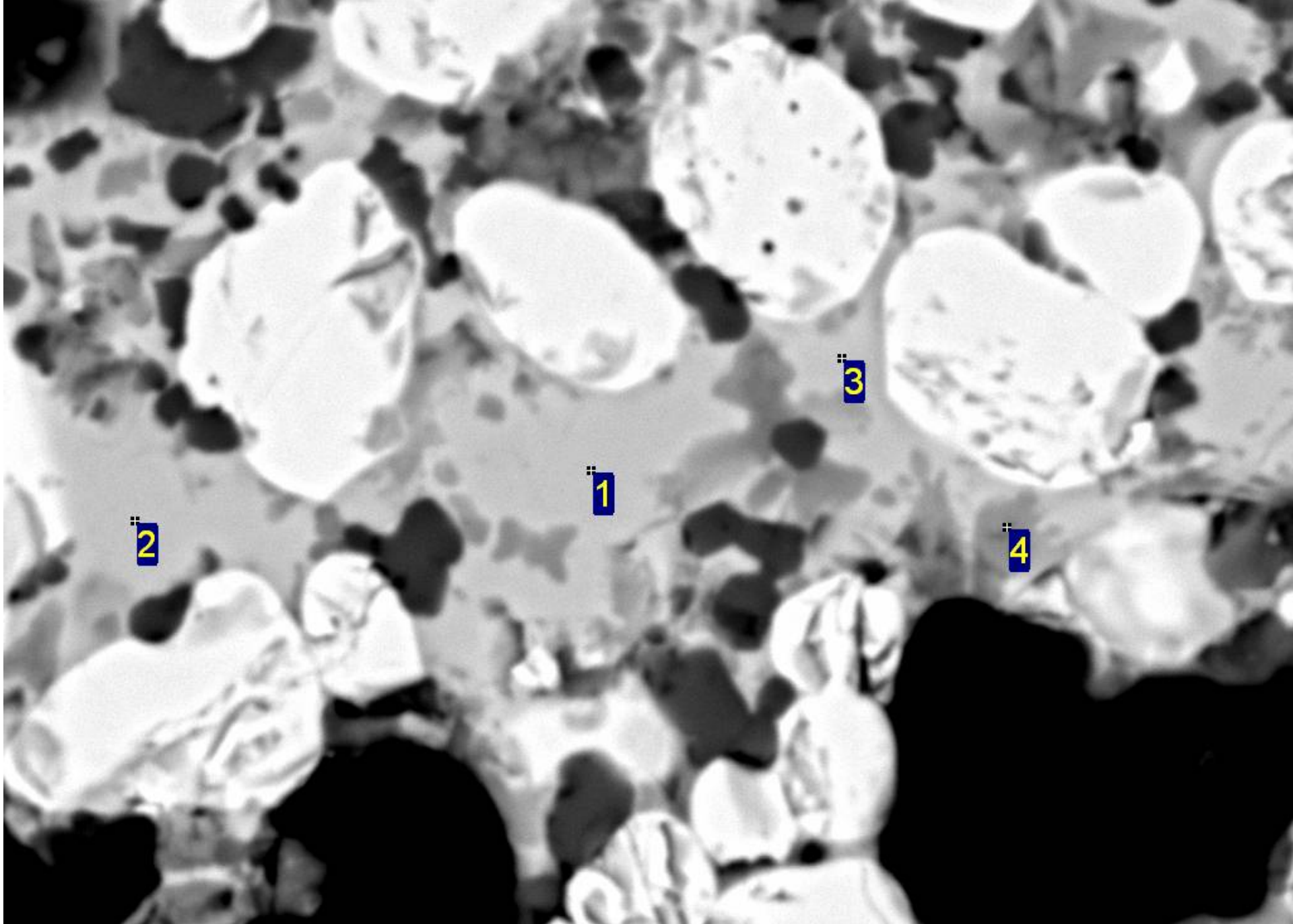


40μm

Electron Image 1

- Evacuated silica tube
- Ash dehydrated at 800 °C
- Hematite-magnetite  $fO_2$  buffer
- Melted at 1081 °C for 6 hours
- Fe-P-O glass between hematite crystals
- Eutectic composition in  $FePO_4$ - $Fe_2O_3$  system

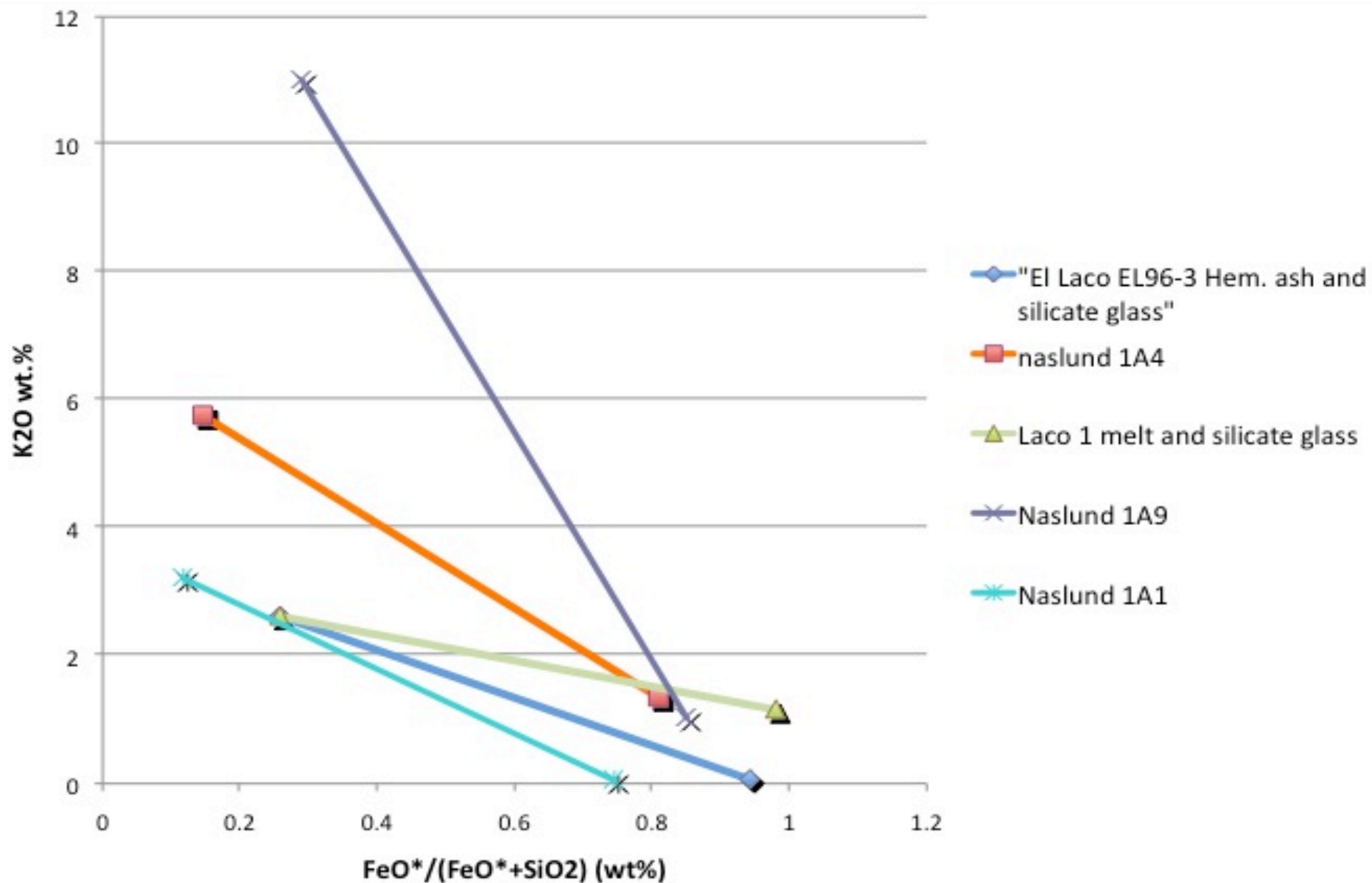




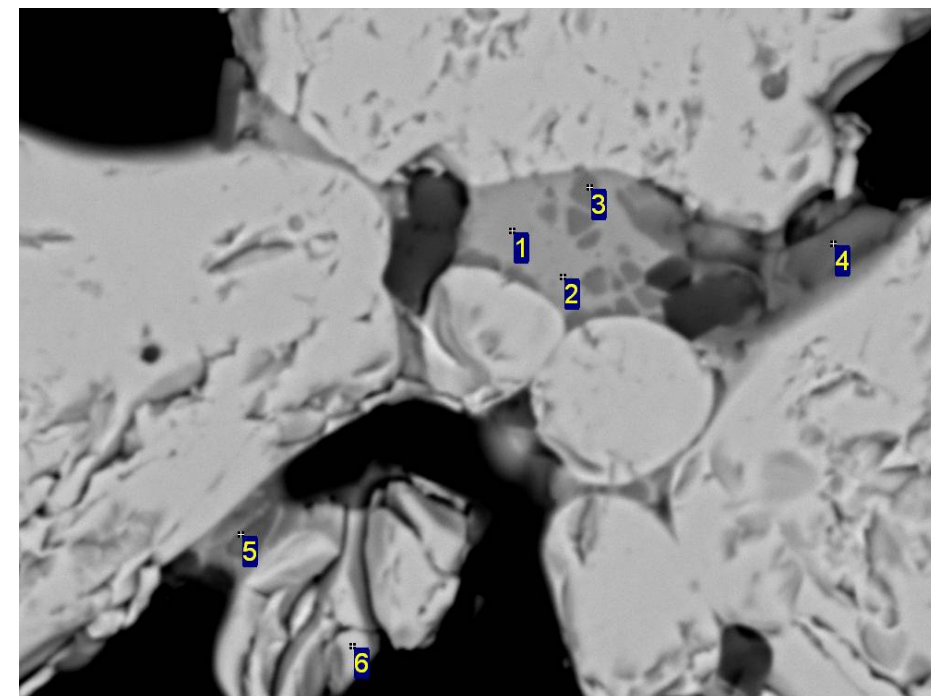
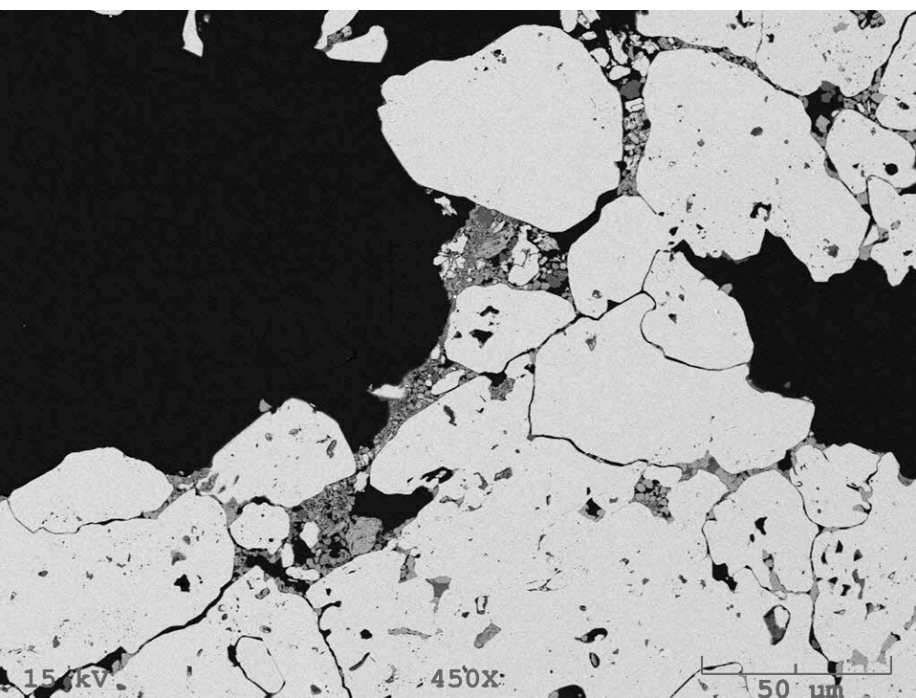
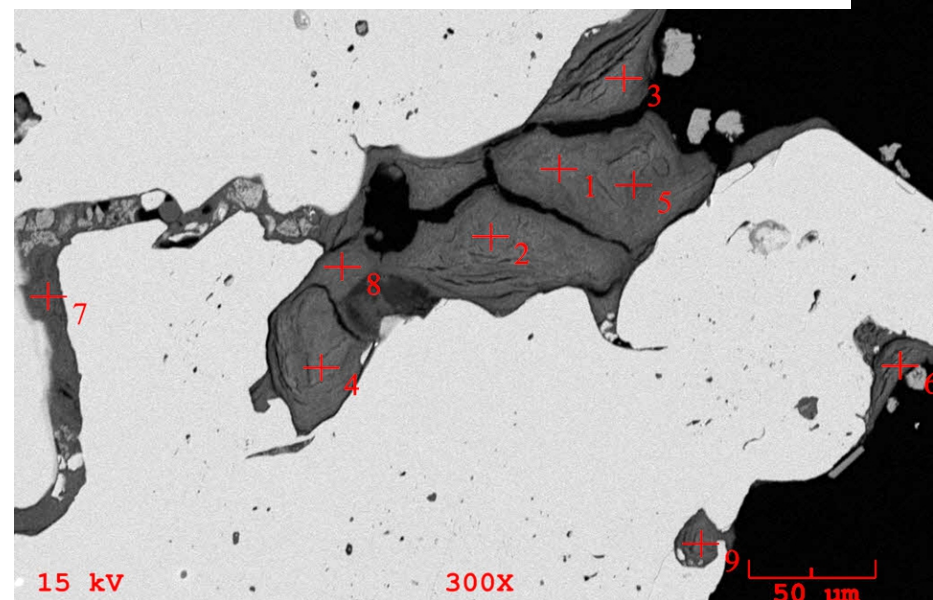
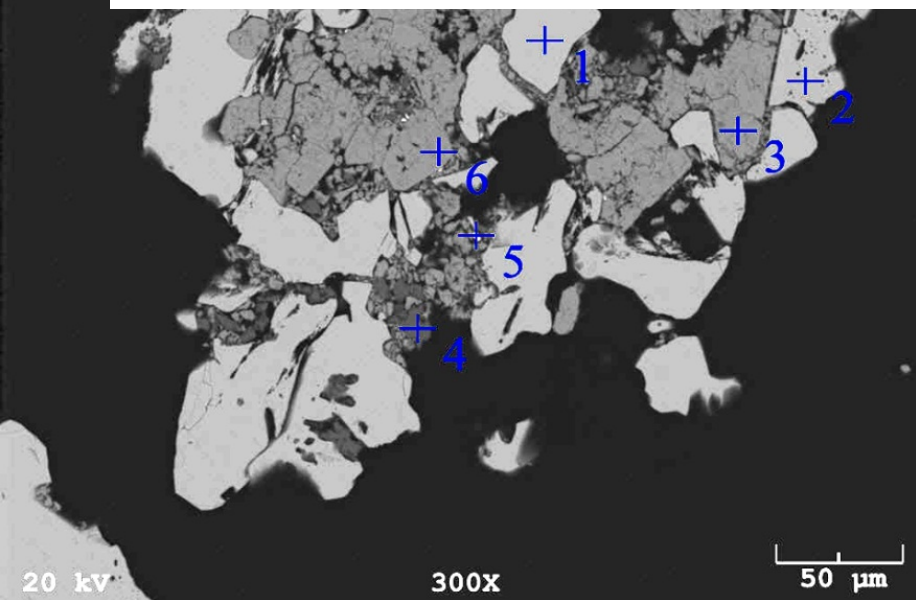
40µm

Electron Image 1

# Shoshonite-ash immiscibility?

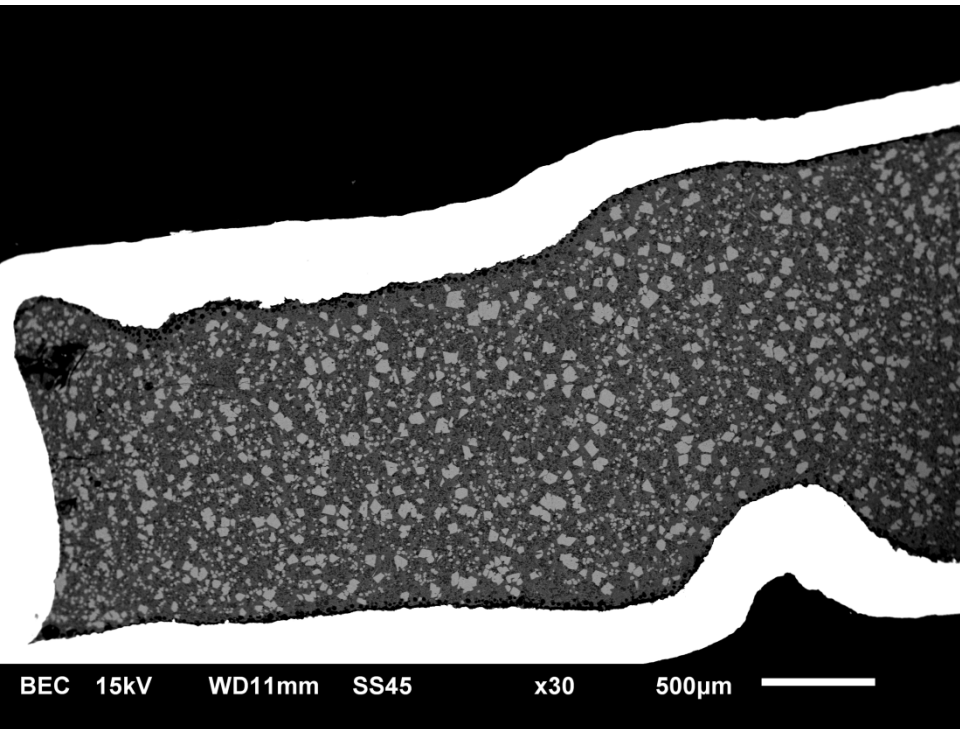


# Textural evidence for two coexisting liquids at El Laco



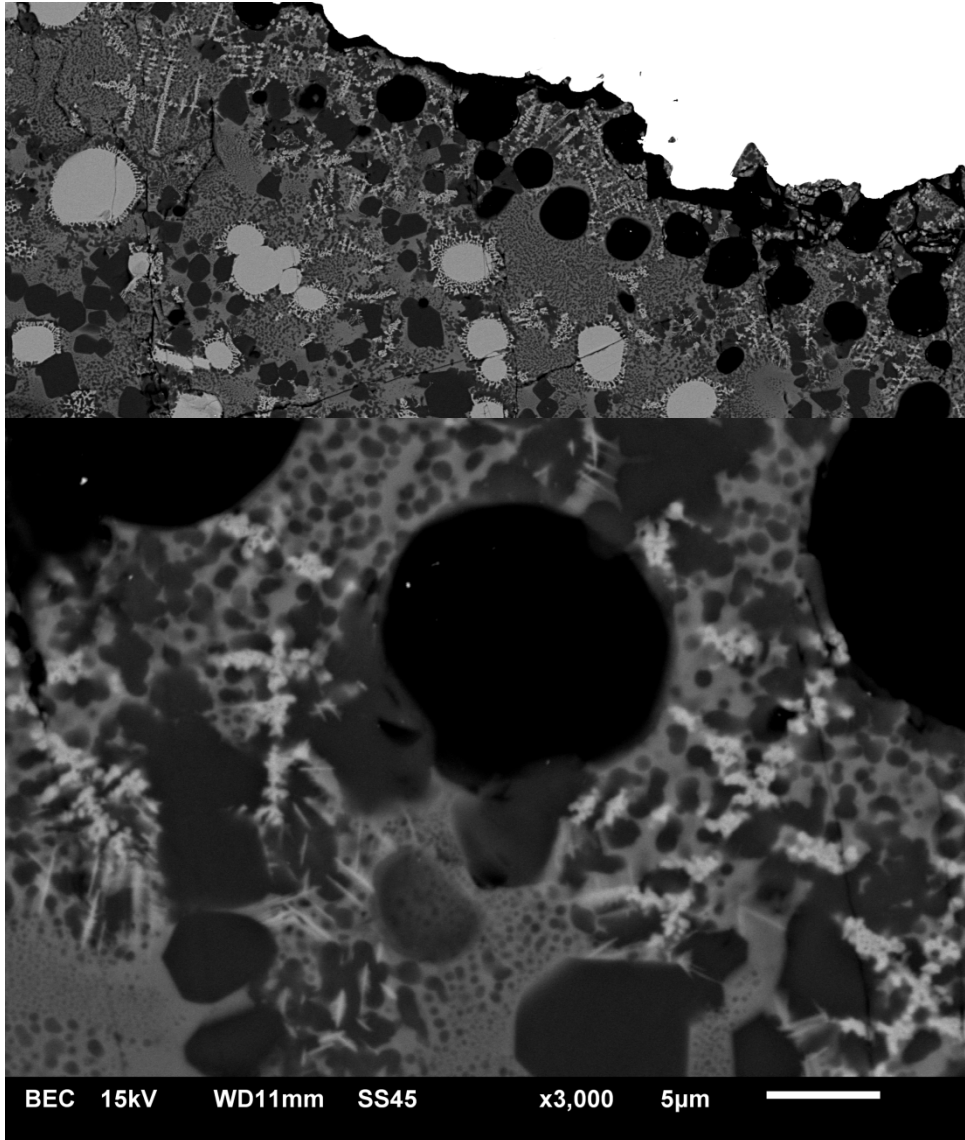


# High pressure melting experiment (U of Toronto)



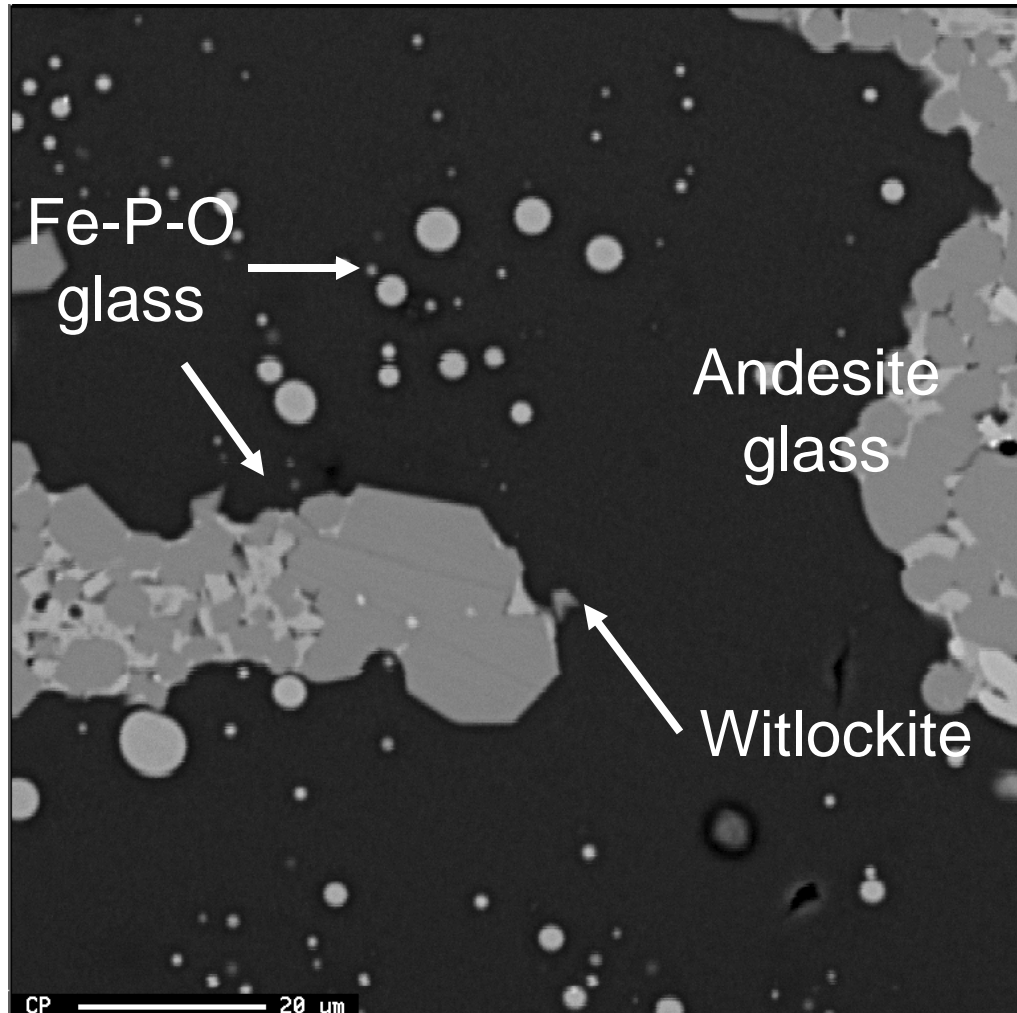
- Synthetic shoshonite glass
- Anhydrous magnetite
- Anhydrous  $\text{FePO}_4$
- Pt capsule
- 1.0 GPa, 1200 °C

# Phase assemblage



- Magnetite
- Hematite
- Silica mineral
- Phosphate-silicate melt ( $P > Si$ )
- Vapour bubbles (possible quench phase)
- No immiscible liquids
- No, the hematite is not an immiscible liquid

# High pressure melting experiment (SUNY)

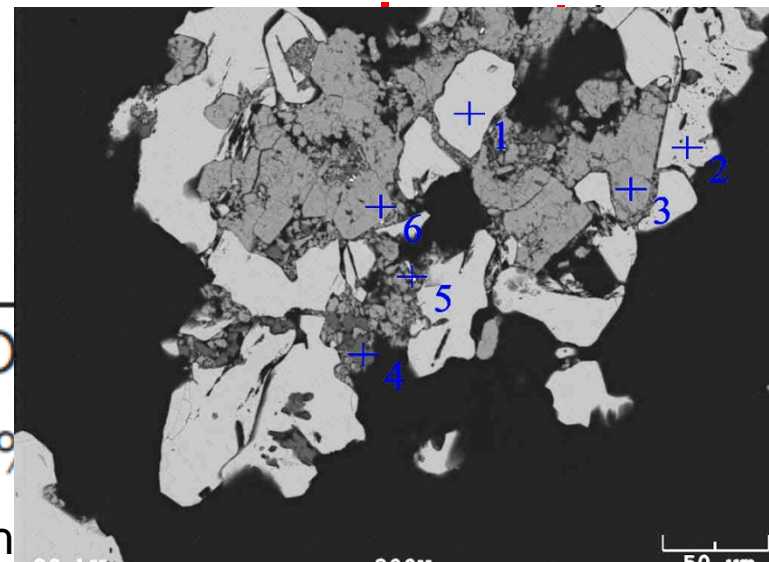
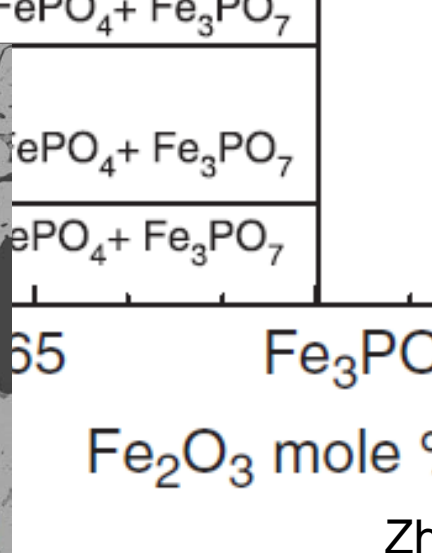
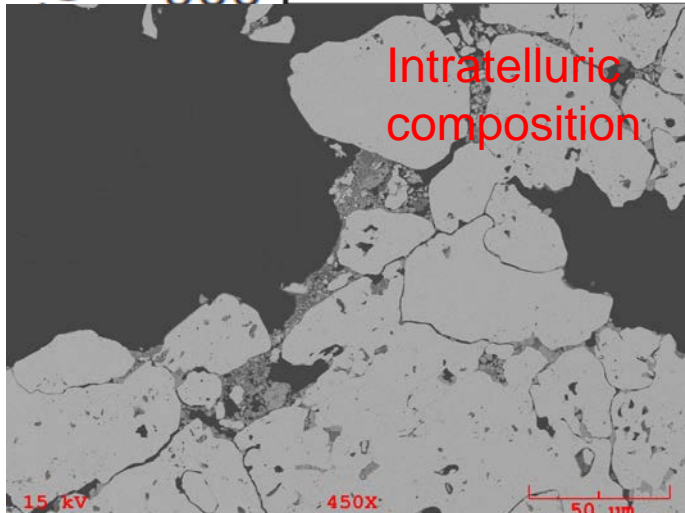
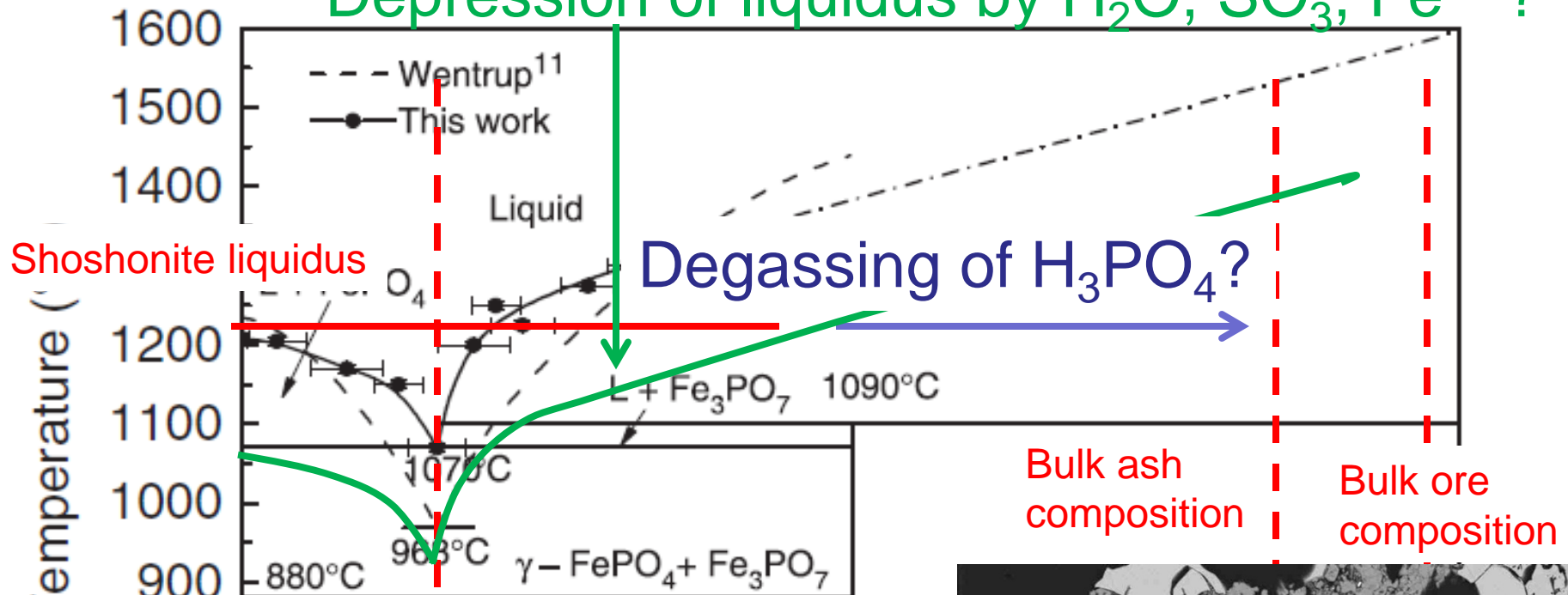


- AGV1 andesite standard
- H<sub>2</sub>O-saturated
- H<sub>3</sub>PO<sub>4</sub> added
- 0.4 GPa, 900 °C
- Immiscible liquids



# Phase relations in $\text{Fe}_2\text{O}_3\text{-FePO}_4$

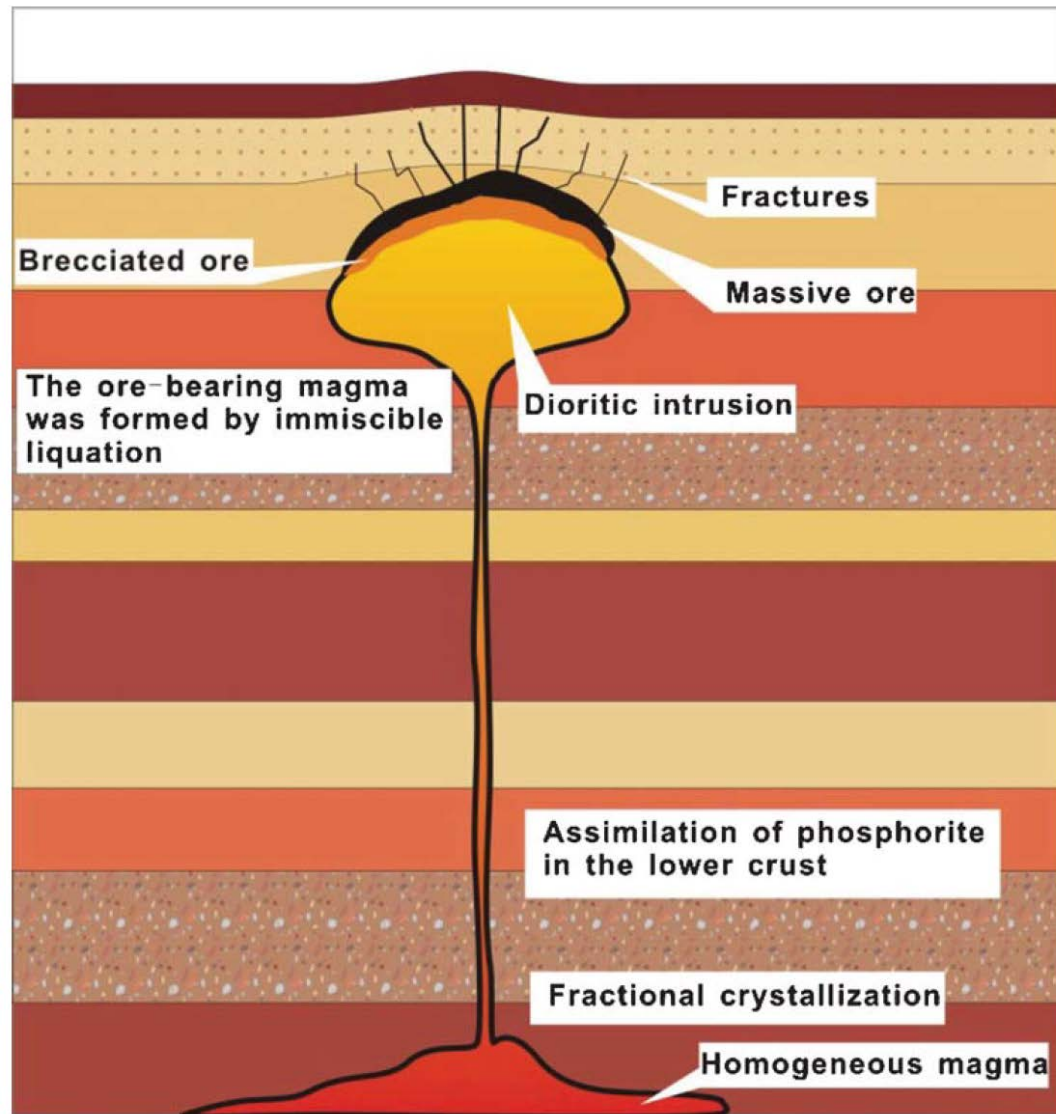
Depression of liquidus by  $\text{H}_2\text{O}$ ,  $\text{SO}_3$ ,  $\text{Fe}^{2+}$  ?



# Summary

- El Lago/Kiruna type Fe ores are magmatic
- Immiscible Fe-rich liquids form in many systems
- FeO-rich liquids are very diverse
- El Lago/Kiruna liquids are unusually low in  $\text{SiO}_2$ ,  $\text{TiO}_2$ , apatite
- The El Lago/Kiruna-type liquid differs from most others in its abundance – Gt compared to milligrams
- The ore magma was close to magnetite composition but must have contained phosphate and other components
- Immiscibility between shoshonite and Fe-P-O liquid
- Additional components likely ( $\text{H}_2\text{O}$ ,  $\text{SO}_3$ , Cl ...)

# Where did the P come from?



- Gushan, China
- Assimilation of phosphorite?
- Liquation of Fe-P-O melt from calcalkaline silicate magma

Hou et al., 2009 Int Geol Rev.



Where did all the P go (how much was there in the first place)?



$\text{H}_3\text{PO}_4?$

# Future directions

- Further experiments – can we reduce the amount of P required?
- Addition of fluxing components  $\text{H}_2\text{O}$ ,  $\text{SO}_3$ , Cl
- Effects of pressure
- Effect of water on liquidus temperature vs liquid-liquid solvus





# Phase relations

$\text{Fe}_2\text{O}_3\text{-P}_2\text{O}_5$

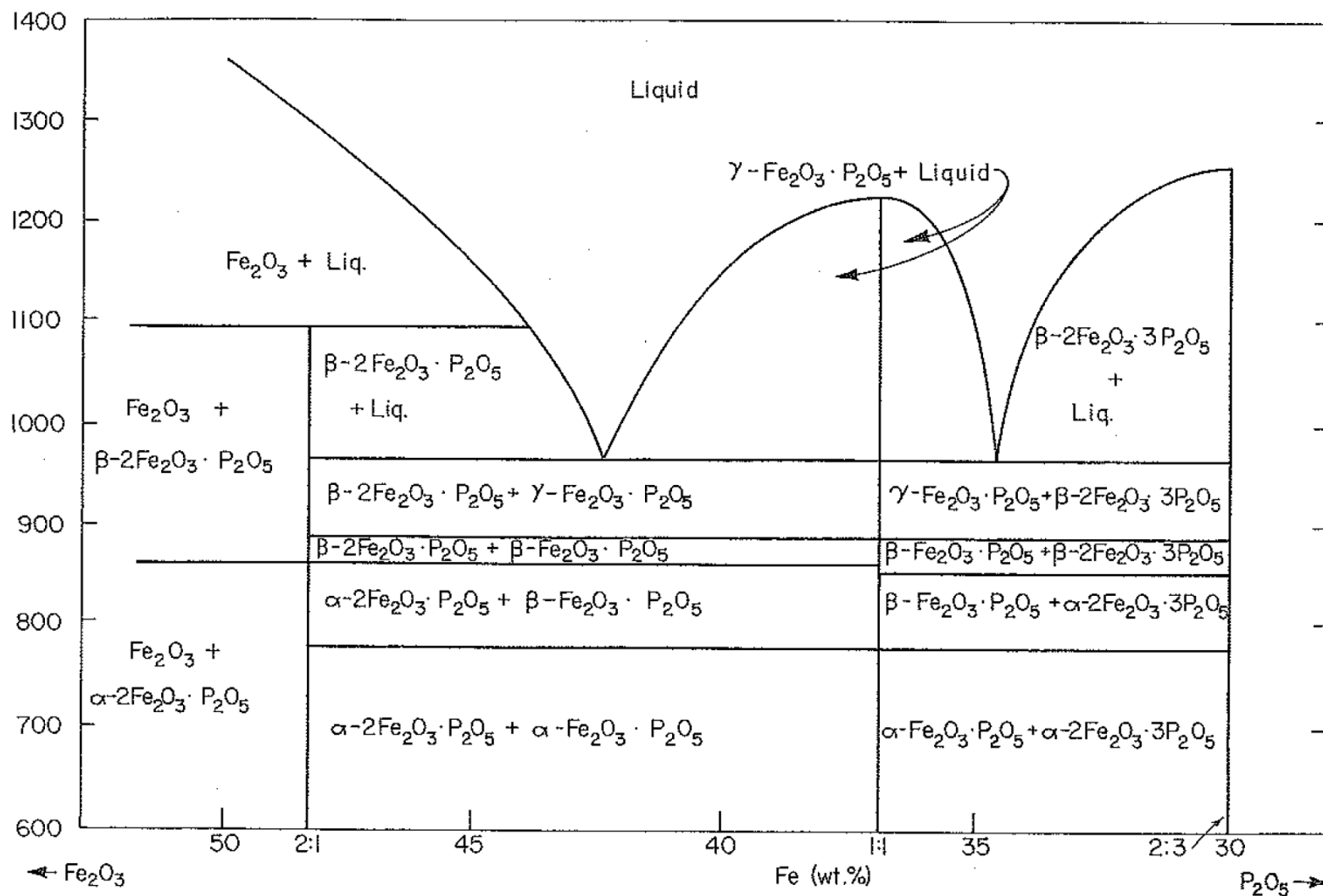
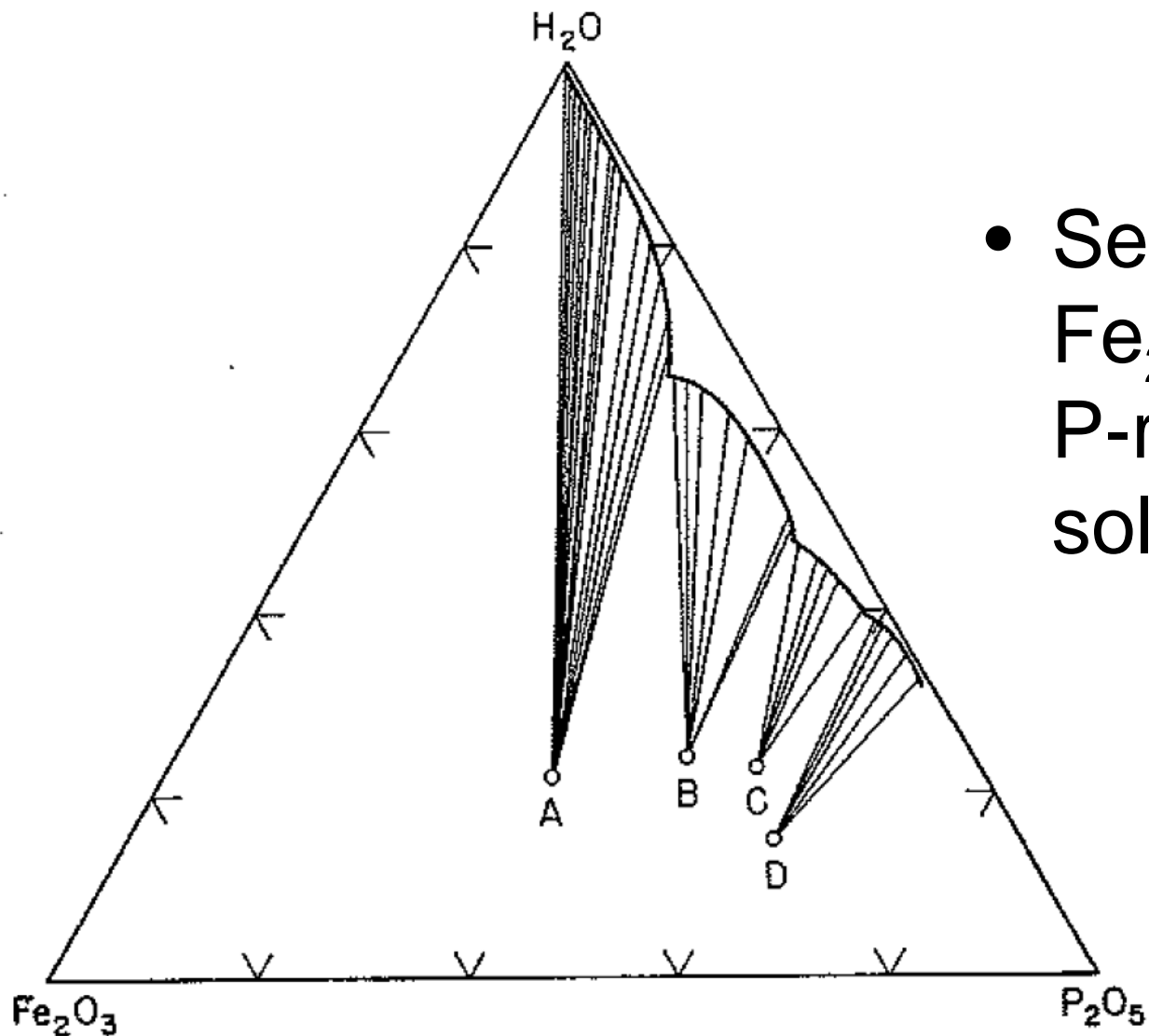


Fig. 338.—System  $2\text{Fe}_2\text{O}_3 \cdot \text{P}_2\text{O}_5\text{-}2\text{Fe}_2\text{O}_3 \cdot 3\text{P}_2\text{O}_5$ .

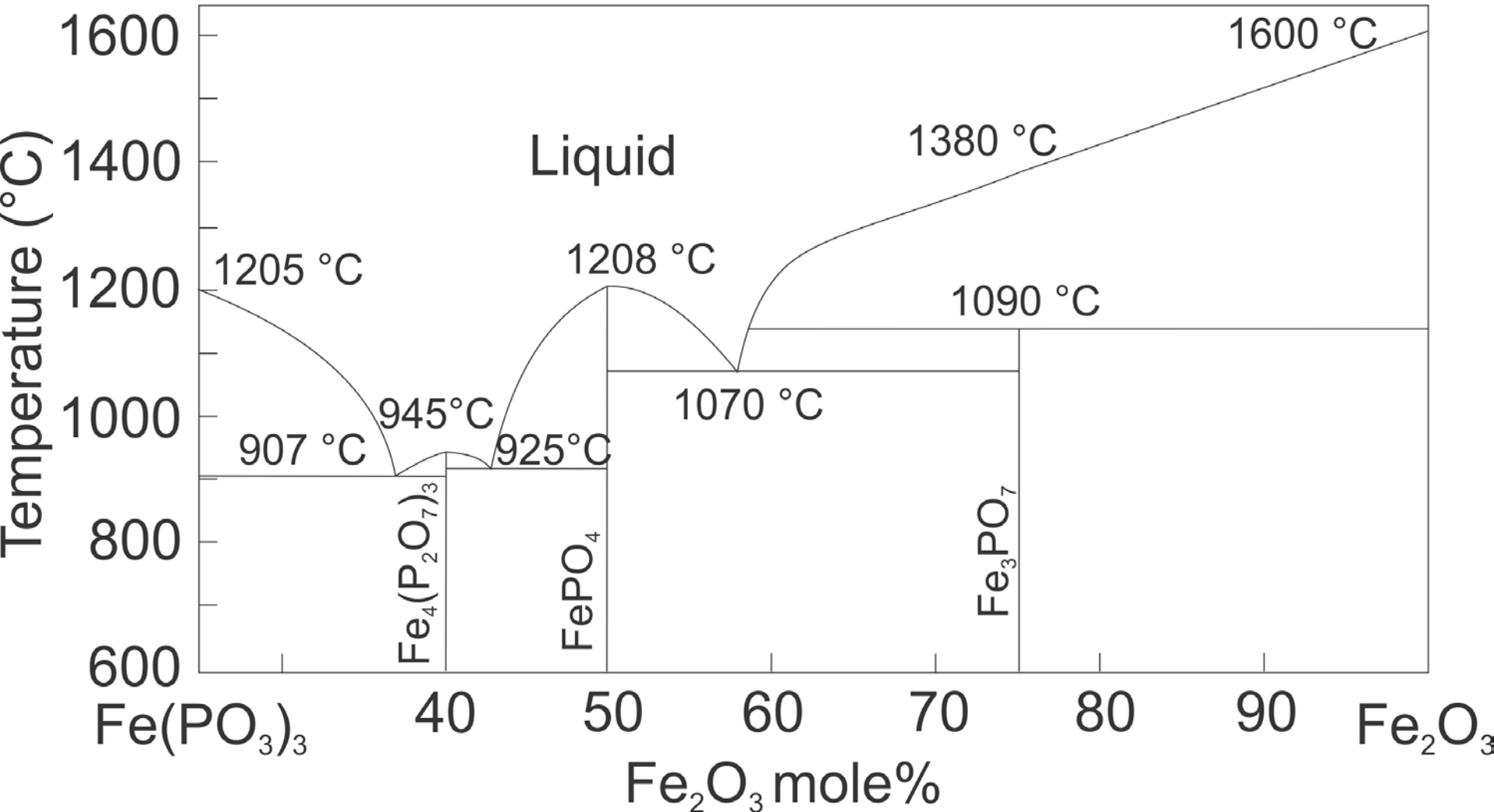
H. Wentrup, *Arch. Eisenhüttenw.*, 9, 57 (1935-36).



- Several wt%  $\text{Fe}_2\text{O}_3$  dissolve in P-rich aqueous solution at  $25^\circ\text{C}$

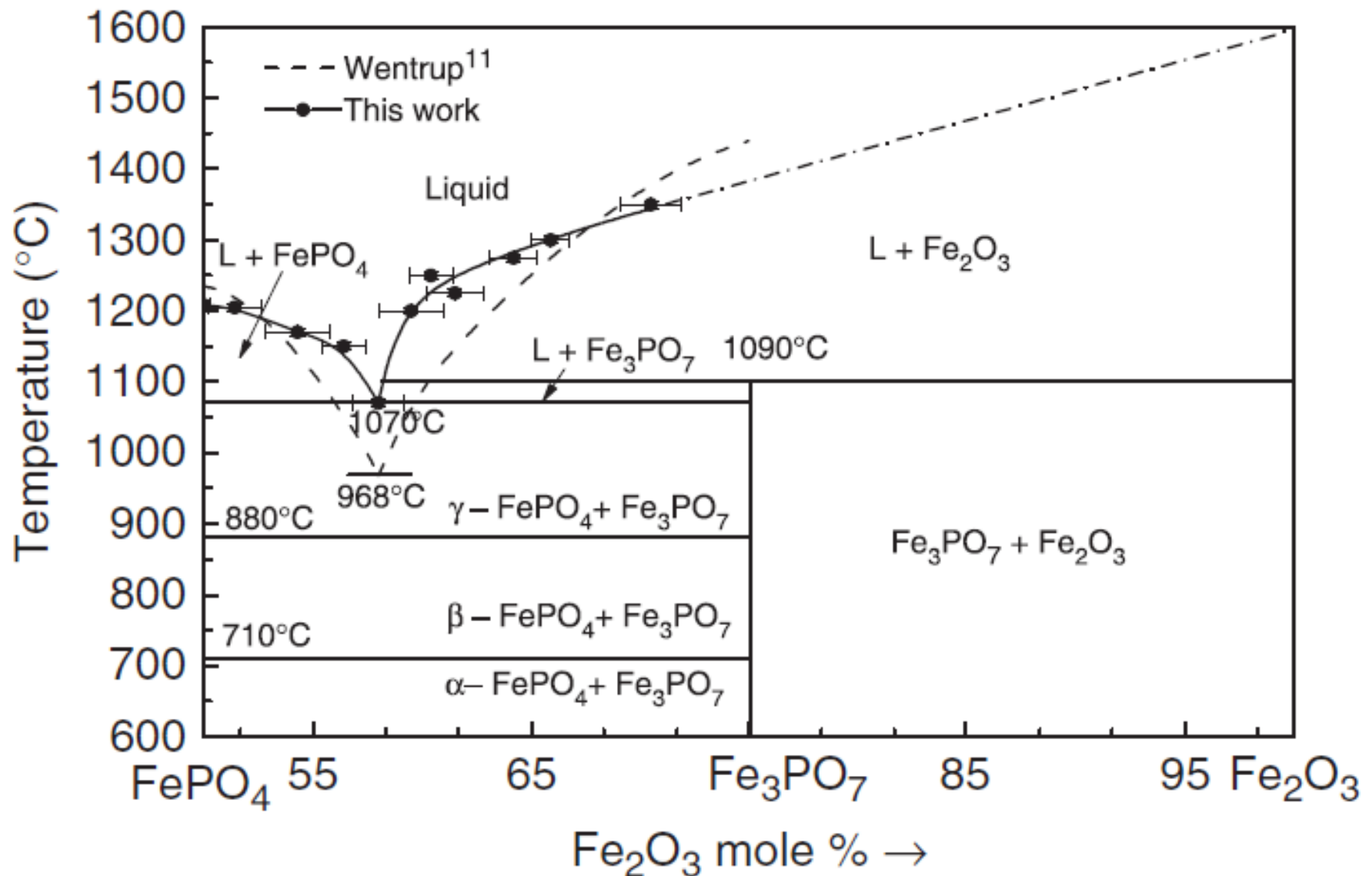
FIG. 2016.—System  $\text{Fe}_2\text{O}_3$ - $\text{P}_2\text{O}_5$ - $\text{H}_2\text{O}$ , at  $25^\circ\text{C}$ . A =  $\text{Fe}_2\text{O}_3 \cdot \text{P}_2\text{O}_6 \cdot 5\text{H}_2\text{O}$ ; B =  $\text{Fe}_2\text{O}_3 \cdot 2\text{P}_2\text{O}_6 \cdot 8\text{H}_2\text{O}$ ; C =  $\text{Fe}_2\text{O}_3 \cdot 3\text{P}_2\text{O}_5 \cdot 10\text{H}_2\text{O}$ ; D =  $\text{Fe}_2\text{O}_3 \cdot 3\text{P}_2\text{O}_5 \cdot 6\text{H}_2\text{O}$ .

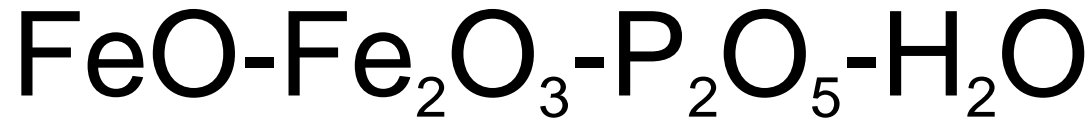
# Iron oxide – phosphate system





# FePO<sub>4</sub>-Fe<sub>2</sub>O<sub>3</sub> (detail)





- Liquids present below 1000°C
- Low viscosity (< 1 PaS)
- Popular choice for nuclear waste storage
- Dissolve abundant oxides, > 1%  $\text{SO}_3$
- $\text{Fe}^{2+}/\text{Fe}^{\text{tot}} > 0.25$  under air, anhydrous
- What happens to eutectic with added degree of freedom in  $\text{FePO}_4\text{-Hm-Mt}$ ?
- At moderate  $f\text{O}_2$ ,  $\text{Fe}_3\text{O}_4$  will likely predominate
- $\text{H}_2\text{O}$  is probably a very effective flux
- $\text{P}_2\text{O}_5$  evaporates rapidly from melts

# $\text{Fe}_3\text{PO}_7$ peritectic at $1090\pm 8^\circ\text{C}$

